Depositional conditions of the coal-bearing Hirka Formation beneath Late Miocene explosive volcanic products in NW central Anatolia, Turkey

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This work focuses on the relationship between the coal deposition and explosive volcanism of the Miocene basin, NW central Anatolia, Turkey. The coal-bearing Hirka Formation was deposited over the Galatian Andesitic Complex and/or massive lagoonal environments during the Miocene. The investigated lignite is a high ash (from 32 to 58%) and sulphur (from 1.43 to 3.03%) lignite which is petrographically characterised by a high huminite content. The mineral matter of the studied lignite samples is made up of mainly clay minerals (illite-smectite and kaolinite), plagioclase and quartz in Bolu coal field, clay minerals (illite-smectite, smectite and illite), quartz, calcite, plagioclase and gypsum in Seben coal field, quartz, K-feldspar, plagioclase and clay minerals (kaolinite and illite) in Kibriscik, and dolomite, quartz, clinoptilolite, opal CT and gypsum in Çamlıdere coal field. The differences in these four types of lignite with specific mineralogical patterns may be due to the explosive volcanic events and depositional conditions which changed from one coal field to the others. There is a zonation from SW to SE in the studied area for zeolites such as Opal CT+smectite-clinoptilolite-analcime-K-feldspar. Carbonate minerals are commonly calcite in Seben and Kibriscik coal fields. In Bolu, coal samples are devoid of calcite and dolomite. These analyses show that there is an increase in the amount of Mg and a decrease in the amount of Na from the northwestern part to the southern part in the study area.

1. Introduction

During the period Neogene to Quaternary, several lacustrine basins developed in Turkey. The Neogene basins were filled by clayey, carbonaceous and sandy sediments, and also by explosive products of contemporaneous K-rich calc-alkaline volcanism with various degrees of crustal contamination (Yılmaz 1989; Güleç 1991; İnci 1991; Gündoğdu *et al* 1996). These basins are generally characterised by volcanic-sedimentary successions. Places including the study area are characterised by important fossil-fuel and industrial-mineral resources such as lignite, oil shale, clays, borates and zeolites (Şener *et al* 1995; Gündoğdu *et al* 1996; Şener and Gündoğdu 1996).

The study area is located between Bolu and Ankara in NW central Anatolia (figure 1a). A large part of the studied Neogene basin is covered by Upper Miocene pyroclastic rocks. The coal-bearing Hirka Formation was first named by Aziz (1976) in the Beypazarı basin, situated approximately 100 km northwest of Ankara in central Anatolia. This formation is understood to have been deposited on the Galatian andesitic complex (or massif) under stable lacustrine conditions. There are four coalfields in the study area: the Bolu coal field, the Seben coal field, the Kibriscik coal field and the Çamlıdere coal field (figure 1b).

The area has been studied from time to time with respect to geological, stratigraphical and volcanological characteristics. The first geological

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Figure 1. Location map of (a) study area and (b) coal seams.

reports of the area were made by Milch (1903); Chaput (1931); Blumenthal (1941); and Türkecan *et al* (1991).

In the present study, I have carried out detailed geological, petrographic and geochemical studies of a large number of samples from all the coal seams measuring ≥ 1 m in thickness. The main objective of this study is to focus on the petrographic composition of coals and other rocks formed on the andesitic complex under stable lacustrine conditions. Further, variations in petrographic composition of the coals and other rocks have been used to reconstruct the palaeodepositional conditions for the coal deposits.

2. Geological setting and tectonic evolution

The coal-bearing strata belong to the Hirka Formation of the Koroğlu Group (table 1). The oldest rocks exposed in the vicinity (Paleozoic) are the Sekli Metamorphics, which comprise the basement

to the Cretaceous sediments. The Cretaceous–Eocene sedimentary sequences of the region are exposed along the south-eastern edge of the Sekli Metamorphics. The lowermost unit of this depositional period is the Keltepe Formation (200 m thick), which rests unconformably on the Sekli Metamorphics and consists of hard, thick-bedded limestone – dolomite of variable colours (grey, white and light yellow) followed by thin beds of softer marl at top.

The Keltepe Formation is overlain conformably by the Gölcük Formation of the Aladağçay Group (150–180 m thick). This formation comprises sandstone, claystone, mudstone, marl and limestone containing Senonian–Maastrichtian foraminifera. The top of the Seben Formation, which consists of marl containing Senonian–Maastrichtian molluscs, referegents the Cretaceous–Tertiary boundary.

The older Tertiary belongs to the Cuisian Emenler Formation that consists of marl (50–75 m thick). The Emenler Formation unconformably overlies all of the older formations (the Gölcük and others). The Taşlık Formation conformably overlies

Table 1. Geological succession of coalfields of investigated area (after Taka *et al* 2001).

Age	Group	Formations	Thickness (m)	Rock types
Quaternary		Alluvium	60	Pebble, sand, travertine
		Unconformity		
Pliocene		Geriş	200	Conglomerate, sandstone, tuff
		Unconformity		
		Young volcanics	100	Basalt
		Uruş	200	Conglomerate, sandstone
Miocene	Köroğlu	Köroğlu volcanics 2	150	Agglomerate, andesite
		Uludere	200–500	Tuff
		Köroğlu volcanics 2	200	Basalt, dasite
		Hirka	150–400	Marl, oilshale, coal
		Unconformity		
Oligocene		Uluköy	250	Sandstone, limestone
Eocene		Akçagil	300	Andesitic lava
		Taşlık	200	Pebblestone, tuff
		Emenler	75	Limestone, marl
		Unconformity		
Upper Cretaceous	Aladağçay	Seben	75–100	Sandstone, marl
		Gölcük	50–80	Sandstone, claystone
Jurassic–Cretaceous		Keltepe	200	Limestone, dolomite
		Unconformity		
Pre-Triassic		Basement		Metamorphic rocks

the Emenler Formation and comprises pebblestone, sandstone, mudstone, tuff and limestone (200 m thick). Nummulitic fauna indicate that the Taşlık Formation is lower–middle Eocene in age. Eocene–Oligocene aged Akçagil Formation consists of interbedded agglomerate, tuff, andesitic lava, sandstone, siltstone, claystone and limestone. The Akçagil Formation is conformably overlain by the Uluköy Formation. The thickness of Uluköy Formation has been found to be quite variable by various investigators; it ranges from a minimum thickness of 250 m up to 2000 m in the Beypazarı basin (Aziz 1976). Fauna and stratigraphic position indicate that the Uluköy formation is of Oligocene age.

The Köroğlu Group is made up of six formations, based on lithological distinctions and fossil content: the Hirka Formation, the Köroğlu Volcanities 1, the Uludere Pyroclastics, the Köroğlu Volcanities 2, the Uruş Formation, and the Young Volcanics (table 1). Miocene sedimentation began with the deposition of the interfingering Hirka Formation and Köroğlu Volcanities 1, Uludere Pyroclastics, which unconformably overlie the Uluköy Formation. The Hirka Formation is coal-bearing and in fact, all of the coal research in the study area has been focused on this unit. The Early Miocene Hirka Formation is 150–400 m thick and comprises of marl, oil shale and coal seams; it

interfingers vertically and laterally with the Uludere Pyroclastics, comprising thick-layered tuff and Köroğlu Volcanities 1, comprising basaltic-dasitic lava. However, generally speaking, the coal-bearing strata of the Hirka Formation lie beneath the Uludere Pyroclastics. Petrographic studies of the tuffs reveal that they are composed of plagioclase, quartz and K-feldspar crystal fragments and lithic fragments in an argillized volcanic-glass matrix. The tuffs are products of extensive volcanic activity, which distributed large amounts of ash over the Kıbrısçık and Seben coal fields, and occur as continuous bands extending 100 km from Çamlıdere to the Bolu coal fields (table 1). According to Taka *et al* (2001), each tuff represents several cycles of eruptive activity and deposition in lacustrine and terrestrial environments by a combination of pyroclastic fall, flow and surge. Also present are tuffites which formed in lakes by transportation of original ash deposits. The Uruş Formation, comprising conglomerate and sandstone, conformably overlies the Hirka, Köroğlu Volcanities 1; Uludere Formations and Köroğlu Volcanities 2 consist of agglomerate and andesite. The thickness of this unit is 200 m and its age is Upper Miocene.

Upper Miocene–Pliocene aged Young volcanics consists of basaltic lava. The Geriş Formation is the youngest unit in the area; it is Pliocene in age and comprises of conglomerate, sandstone and tuff;

Figure 2. Geological map of the study area.

Figure 3. Measured stratigraphic section of coal seams.

200 m thick, this formation unconformably overlies all other lithostratigraphic units.

The tectonic evolution of Turkey since the early Miocene has been controlled by two collisions along the İzmir–Ankara–Erzincan and Bitlis sutures (figure 1a). N–S intracontinental convergence associated with the Palaeocene–Eocene İzmir–Ankara–Erzincan suture continued until the late Miocene (Şengör and Yılmaz 1981). On the other hand, the Middle Miocene Arabia–Eurasia collision, along the Bitlis suture in south-eastern Turkey (Şengör and Yılmaz 1981), accepted as the beginning of the neotectonic period in Turkey (Şengör 1980), has generated a N–S extensional regime in western Turkey since the late Miocene (Şengör *et al* 1985).

During this tectonic evolution, a large number of northeast- and northwest-trending basins have developed within the İzmir–Ankara suture zone, which is occupied by ophiolitic, metamorphic and sedimentary rocks of Permian to Eocene age. In the basin, the main tectonic structures affecting the Miocene units are oblique faults and roll-over type folds, which are generally parallel to the trends of the basins and the İzmir–Ankara suture.

3. Mode of occurrence of the miocene coal seams

3.1 Coal fields of Bolu

The coal fields of Bolu are the Düvenlik and Çaygökpınar fields. There are two coal seams in

these fields. The upper seam is about 0.10–0.25 m thick while the lower seam is about 0.10–0.45 m thick (figures 2 and 3).

3.2 Coal fields of Seben

The major coal fields of Seben are the Hoçaş and Kaşbıyıklar. The coal deposit is situated in an approximately 4 km-long, E–W-trending horizon. There are three coal seams in the Hoçaş coal field. The upper seam is about 0.50 m thick, the middle seam is 0.25 cm thick, while the lower seam is 0.50 m thick. All of these seams are still being mined actively. A 0.10 m-thick claystone bed is present between the middle and lower seams. There are two outcrops in the Kaşbıyıklar coal field: one of them has one seam 0.20 m thick, while the other has three coal seams that range from 0.10 to 0.25 m in thickness. There is a 0.20 m-thick tuff layer between the middle and upper seams in the Kaşbıyıklar coal field (figures 2 and 3).

3.3 Coal fields of Kırıbrıcık

The coal fields of Kırıbrıcık are the Kabalar and Bahçeköy fields. The Kabalar coal field is located north of Beypazarı at a distance of about 45 km on the Beypazarı–Ankara road. There is one 0.80 m-thick coal seam in the Kabalar field, while the Bahçeköy coal field has two coal seams. The lower seam is 0.80 m thick while the upper seam is 0.10 m thick; the two seams are separated by thick tuff and limestone interlayers (figures 2 and 3).

Table 2. Results of proximate and ultimate analyses.

	Bolu	Seben	Kıbrısçık	Çamlıdere
Moisture %	30	32	25	29
Ash %	32	36	38	37
Volatile matter %	14	10	9	5
Fixed carbon %	24	22	18	19
Nett calorific value kcal/kg	4376	3460	1185	3763
Gross calorific value kcal/kg	4875	3852	1362	3995
Combustible S %	1.46	0.01	1.64	2.01
Sulphur in ash %	0.64	1.42	0.06	1.02
Total S %	2.10	1.43	1.70	3.03
% Rmax	0.61	0.36	0.34	0.37

3.4 Coal fields of Çamlıdere

The minor coalfields of Çamlıdere are the Buğralar and Elmalı fields. There are three coal seams in the Buğralar coal field. The upper seam is about 0.15 m thick and the middle seam is 0.10 m thick, while the lower seam is about 0.65 m thick. The Elmalı coal field has three coal seams whose thicknesses vary from 0.05 m to 0.20 m (figures 2 and 3).

4. Methods of study

Pillar coal samples were taken at a series of outcrops and in galleries; these samples reflected the exposed thickness of the seam at the sampling point (figures 2 and 3). Proximate and ultimate analyses as well as petrographic analyses were carried out. Additionally HI, OI, TOC and mineralogical contents were determined.

For coal-petrographic analyses, samples were crushed to less than 1 mm. Maceral analyses were conducted on polished pellets using 32 × oil-immersion objectives. Fluorescent light was also used to identify the liptinite macerals. Petrographic analyses were based on point-counting methods; for maceral analyses, 250–300 points were counted, whereas 100 readings were taken for reflectance measurements. Maximum reflectance measurements on eu-ulminites A and B were carried out using a reflected-light microscope (Leitz MPV SP) with a 3 × oil-immersion objective and sapphire standard for calibration (0.551% R).

The contents of total organic carbon (TOC) and total sulphur (TS) were carried out on samples with a LECO-CS analyser; carbonate was removed using dilute hydrochloric acid. Calorific values were determined using calorimeters after BIN 51.900. The mineralogical compositions of 58 samples were determined via XRD analysis.

5. Results and discussion

5.1 Proximate and ultimate analyses

Table 2 shows the results of the proximate and ultimate analyses of composite profile samples from the four coal fields in the study area. Ash contents range from 32 to 58%. The Kıbrısçık sample is characterised by the highest ash content and the lowest calorific value, whereas the Bolu sample has the lowest ash content. The higher calorific values range from 1362 to 4857 kcal/kg.

Total sulphur contents range from 1.43 to 3.03%. The Çamlıdere lignite has the highest S total content (3.03%), whereas the Bolu and Seben coal fields have intermediate levels (1.70–2.10%), and the Seben coal fields has lower S total (1.43%). The combustible S content of the Kıbrısçık sample is higher than the samples from the other coal fields, and the pyritic S contents of the Seben and Çamlıdere coal-field samples are higher than those from other fields. These results indicate that organic sulphur contents in the study area decreased from SW to S where tuff deposition was greater.

5.2 Rank

Volatile-matter contents and vitrinite reflectance have been used as parameters for determining the approximate rank of the coals. The volatile-matter contents (d.a.f. basis) of samples from the studied coal fields range from 24 to 44% (table 2). The vitrinite reflectance (Ro mean) of samples from the studied fields ranges from 0.25 to 0.83 (figure 4).

The Bolu coal samples have volatile-matter contents and Rm% higher than the samples from the other coal fields. The volatile-matter contents (24–44%) and Ro% mean values (0.35–0.61) suggest that these coals range in rank from

Figure 4. Coal-rank classification based on ASTM D-388-1984 and Teichmuller and Teichmuller (1982). **Maf**: moist ash-free basis; **daf**: dry ash-free basis. **M**: Muğla basin samples (from Querol *et al* 1999); **H**: Himmetoğlu samples (from Şener and Gündoğdu 1996); **Be**: Beypazarı samples (from Whateley and Tuncalı 1995); **So**: Soma basin samples (from Şengüler *et al* 2000); **I**: studied samples.

Table 3. Mean percentages of macerals for coal samples from the investigated area.

Macerals			Bolu	Seben	Kıbrısçık	Çamlıdere		
Humunite	{	Humotelinite	{	Textinite	3	2	3	2
				Texto-ulmunite	6	4	3	5
				Eu-ulmunite	8	5	2	4
	{	Humodetrinite	{	Attrinite	2	3	0	4
				Densinite	4	5	4	4
	{	Humocolinite	{	Gelinite	35	34	8	32
			Corpohumunite	0	1	0	1	
Total humunite			58	54	20	52		
Liptinite	{	Sporinite	3	3	2	2		
		Alginite	0	1	0	0		
		Resinite	0	0	0	1		
		Cutinite	2	1	2	2		
		Liptodetrinite	2	1	1	1		
Total liptinite			7	6	5	6		
Inertinite	{	Fusunite	0	0	1	0		
		Macrinite	2	1	2	1		
		Sclerotinite	0	0	0	0		
		Liptodetrinite	1	1	1	1		
Total inertinite								
Pyrite		8	5	6	5			
Clay minerals and silica		24	33	55	35

Figure 5. Ternary diagram showing the maceral distribution in composite coal seams of the study area (B: Bolu, S: Seben, K: Kırıscık, Ç: Çamlıdere).

approximately lignite to high volatile “C” bituminous (figure 4), as per German (DIN) and North American (ASTM) classifications.

In some samples, calorific values, volatile-matter contents, elemental carbon and Rm coincide in this slightly higher rank coal, such as in the Kırıscık lignites. This higher degree of maturation is probably attributable to more extensive tuff deposition and lava interlayers in these particular regions, as compared to the other coal fields.

Maceral distributions of the composite samples are given in table 3. As seen in this table and in figure 5, the lignite samples are characterised by high huminite contents, mainly gelinite macerals. The lignites do not exhibit cellular structures but are mostly gelified forms of huminite. Contents of lipitine macerals (mainly sporinite and cutinite) are slightly higher than the inertinite macerals (mainly macrinite). The lignite is also characterised by a high mineral-matter content, mainly clay minerals, quartz, calcite and pyrite.

The tissue preservation index (TPI) and gelification index (GI) were calculated according to the methodology of Toprak (1996). Figure 6 gives a TPI/GI cross-correlation plot. These results indicate a limnic depositional environment for the studied lignite.

5.3 Mineralogy

5.3.1 Bolu coal fields

The mineral matter of the Düvenlik and Çaygökpınar coal samples comprises mainly clay minerals (illite-smectite and kaolinite), plagioclase and opal CT (table 4).

5.3.2 Seben coal fields

The mineral composition of the Hoçaş coal samples consists mainly of clay minerals (illite-smectite,

Figure 6. Facies diagram (TPI vs. GI) and suggested depositional environments for the studied coals (B: Bolu, S: Seben, K: Kırıscık, Ç: Çamlıdere) modified from Toprak (1996).

smectite, illite), quartz, calcite, plagioclase and, locally, gypsum. The Kaşbiyıklar coal samples contain clay minerals (illite-smectite, smectite, illite), calcite, plagioclase and quartz, whereas in samples K-2 and K-3, K-feldspar is present and in sample K-2 opal CT and clinoptilolite occur (table 4).

5.3.3 Kırıscık coal fields

The mineral matter of Kırıscık coal samples comprises mainly quartz, K-feldspar, plagioclase and clay minerals (kaolinite, illite). In addition to these minerals, analcime (in Ka-4 and Ka-6), gypsum (in Ka-5), and dolomite (in Ka-1 and Ka-6) were identified (table 4).

5.3.4 Çamlıdere coal fields

The Çamlıdere coal samples contain dolomite, quartz, clinoptilolite, opal CT and gypsum (table 4).

Bearing in mind the ash and mineral-matter contents of the studied samples, two different lignites can be distinguished: (i) the Kırıscık and Çamlıdere lignites, which are characterised by high-medium ash contents (38–58%) and relatively high proportions of detrital minerals, which were derived from explosive volcanic events; (ii) the other lignites, which have relatively high proportions of clay minerals with variable ash contents (32–36%). These two different types of lignite, with different specific mineralogical patterns, may be a result of explosive volcanic events and resultant changing depositional conditions from one coal field to another. There is a general trend of decreasing kaolinite contents from Bolu (NW part

Table 4. Mineral distributions of coal samples from the investigated area.

Location	Sample	XRD															
		I	S	I-S	K	Q	C	P	F	Kl	J	CT	A	Pi	Ar	Do	An
Düvenlik	1			1	2							3	4				
Düvenlik	2		1		2			3									
Hoçaş	1		1					2									
Hoçaş	2	2		1		4		5				3	6				
Hoçaş	3			3		4	1	5				2					
Hoçaş	4a		3	2		5	1	6				4					
Hoçaş	4b		3			4	2	1			5			6			
Hoçaş	5	2	1			4	6	3			5						
Kaşbıyıklar	Köm 1	6		1		4	2	5	6			3					
Kaşbıyıklar	Köm 1a		3	4			1	2			5						
Kaşbıyıklar	Köm 1b		1			3		2		4	5						
Kaşbıyıklar	Köm 1c		2				1	3									
Kaşbıyıklar	Köm 1d		1				3	2				4					
Kaşbıyıklar	Köm 1e	4	1				3	2									
Kaşbıyıklar	Köm 1f	5	1				3	2							4		
Kaşbıyıklar	Köm 2	5		1			2	3	4	7		6					
Kaşbıyıklar	Köm 3			1		3	2	4	5				6				
Kaşbıyıklar	Köm 4	4		3		5	2	1									
Kaşbıyıklar	Köm 5	3	1	2		4		5									
Kaşbıyıklar	K 1			1		5	3	4				2					
Kaşbıyıklar	K 2	7		1		3	2	4	7	6		5					
Kaşbıyıklar	K 3			1		4	2	5	6				3				
Kaşbıyıklar	K 4	4		3		5	2	1									
Kaşbıyıklar	K 5	2	1		3	4											
Kabalar	Ka 2				4	1		3	4								
Kabalar	Ka 3	5		4	3	1			2								
Kabalar	Ka 4	7		6	2	1		3	4								
Kabalar	Ka 5				5	1	4	2	3		6						
Kabalar	Ka 6	7	6		5	4			3							2	1
Kabalar	Ka 7	4	1			3		2									
Buğralar	Bu 26					2				3		5	4			1	

I: Illite; **S:** Smectite; **I-S:** Illite-Smectite; **K:** Kaolinite; **C:** Calcite; **P:** Plagioclase; **F:** K-Feldspar; **J:** Gypsum; **CT:** Opal CT; **A:** Amorph; **Pi:** Pyrite; **Do:** Dolomite; **An:** Analcime; **Q:** Quartz; **Ar:** Aragonite and **Kl:** Clinoptilolite.

of the region) to Seben (S part of the region). Coal deposition in this region was syngenetic with volcanic-ash deposition.

Gypsum in the Seben coal samples indicate a low water table in the environment due to evaporation; thus the depositional environment was quite shallow and the water table was beneath the peatland. Coal samples taken from the southern part of the study area (Kıbrısık, Seben and Çamlıdere) contain opal CT, clinoptilolite and analcime. Volcanic glass, derived from the Uludere Pyroclastics, was probably the major detrital (mineral matter) component in the precursor peat bogs of the studied coal fields. The presence of different zeolite

minerals indicates that there probably were different proportions of Na and/or K to Ca, and Al to Si, in the precursor volcanic ash or in the pore fluids. Analcime crystallization reflects a high Na:Ca ratio in the activation solution and relatively high Al:Si ratio in the volcanic ash, whereas clinoptilolite reflects somewhat lower Na/K:Ca and Al:Si ratios (Gündoğdu *et al* 1996; Temel and Gündoğdu 1996; Whateley *et al* 1996). From SW to SE in the study area, there is a zonation as follows: opal CT + smectite→clinoptilolite→analcime→K-feldspar. Carbonate minerals are relatively abundant in the coals of the Seben and Kıbrısık fields, whereas the Bolu coal samples lack calcite and dolomite. From

NW to S in the study area, there is increasing Mg and decreasing Na. In particular, the Çamlıdere coal samples have higher dolomite contents than do samples from the other coal fields; thus its depositional environment was probably relatively deeper than the others.

6. Conclusions

The coal-bearing Hirka Formation in NW central Anatolia formed under initially favourable, warm, humid, tropical-subtropical palaeoclimatic conditions which later evolved into slightly drier and cooler conditions. The main features of these lignites resulted from the tectonic evolution and volcanic features of the basin due to the influence of tectonics and volcanism on the sedimentation and subsidence rates and on basin morphology.

According to previous research on the geochemical conditions of the Mesozoic and Tertiary coal deposits of the Mediterranean region, all coal seams were deposited in alkaline environments, with high bacterial activity and high organic sulphur contents (Querol *et al* 1999). Not only the geochemical features but also the mineralogical and sedimentological features show that volcanic activity had a major influence on fossil-fuel deposition (namely, coal and oil shale) especially in Turkey during the Neogene epoch and/or neotectonic period.

A key result of this study of the Hirka Formation is recognition of the major influence exerted on the early diagenetic evolution of the coal by explosive volcanic events proximal to the lacustrine environments and has very different properties than the other coal fields of Turkey (figure 5). To date, the thick pyroclastic cover in the ~1100 km² study area has had a profoundly negative effect on coal exploration. However, exploratory drilling – guided by the results of the present study – may lead to the discovery of major new coal deposits.

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