Cosmogenic isotope and track records in meteorites and their implications to cosmic ray fluxes

NARENDRA BHANDARI

Physical Research Laboratory, Navrangpura, Ahmedabad 380 009, India

MS received 15 March 1985; revised 23 April 1986

Abstract. Study of several cosmic ray effects, such as VH track density, spallogenic ^{26}Al and ^{53}Mn activity, ^{21}Ne and $^{22}\text{Ne}/^{21}\text{Ne}$ ratio, made in the same sample or in cores taken from different meteorites can identify parameters related to the exposure history of meteorites and cosmic ray flux variations. Meteorites with single or multiple exposure can be distinguished from a track production rate – $^{22}\text{Ne}/^{21}\text{Ne}$ correlation diagram and cosmic ray flux variations over 10^6 – 10^7 years can be deduced from a three-isotope correlation diagram of ^{26}Al , ^{53}Mn and ^{21}Ne . Isotopic data based on chondrites with simple, one-stage exposure are consistent with the same average galactic cosmic ray intensity over the past 2 million years as that during the past 10^7 years.

Keywords. Cosmogenic isotopes; meteorites; cosmic ray fluxes.

1. Introduction

The possibility that meteorites, having long term cosmic ray records in the form of isotopic products resulting from nuclear reactions and heavy nuclei (VH) tracks, can be used for monitoring cosmic ray fluxes in the past has been discussed since the early work of Arnold *et al* (1961). Calculations show that the rates of production of these cosmic ray effects depend significantly on the shielding depth within a meteorite (Bhattacharya *et al* 1980; Eberhardt *et al* 1966; Fleischer *et al* 1967). Also, the meteorites are not always exposed to cosmic rays in a fixed geometry because of multiple exposure on their parent bodies (Bhandari and Potdar 1982), fragmentation in space and space erosion. The effects of these processes have to be quantitatively assessed before the isotopic data in meteorites can be used to infer cosmic ray intensity. Because of these complications, there is considerable debate on the changes in the cosmic ray flux inferred from meteorite data. Nishiizumi *et al* (1980) and Müller *et al* (1981) have proposed that the cosmic ray intensity over the past 2×10^6 years was higher as compared to the average intensity before that. This is based on statistical analyses of ^{26}Al , ^{53}Mn and ^{21}Ne data after making empirical corrections for shielding depths but without considering the effects of various processes cited above. Recently Moniot *et al* (1983) have measured ^{10}Be , reanalysed the data on several isotopes and concluded that temporal variations in cosmic ray flux have not occurred.

Here we discuss isotopic and track density data in some chondrites restricting ourselves to measurements which have been made in the same sample. Whereas the track density enables us to determine the shielding depth in the case of the same sample measurements, errors in shielding corrections become negligible since the shape of the production profiles of these isotopes are nearly identical. Further excluding meteorites

Table 1. Track density and isotopic data in chondrites.

Meteorite (class)	Pre-atmos- pheric radius (cm)	Track density $\equiv P$ ($\times 10^4/\text{cm}^2$) (A)	^{26}Al (dpm/kg) (B)	^{53}Mn (dpm/kgFe) (C)	^{21}Ne (10^{-8} ccSTP/g) (u) (D)	$^{22}\text{Ne}/^{21}\text{Ne}$ (u) (E)	Exposure age (M.Y.) (u)	Remarks (v)
Bansur (L6)	15 (e)	150 (e)	48 $\pm$ 5(h)	370 (e)	5.35	1.14	15	(A, B, C, D, E)
Barwell (L5)	17 (c)	104 $\pm$ 7(c)	63 $\pm$ 9(g)	203 $\pm$ 23(g)	1.18(g)	1.12	2.5	(A, B, C, D, E)
Bruderheim (L6)		25 $\pm$ 5.5(b)	78 $\pm$ 2(g)	466 $\pm$ 42(g)	11.5(g)	1.10(i)	23.2	(A, B, C, D, E)
Dalgaty		< 5 (h)	52 $\pm$ 2.9(s)	247 $\pm$ 20(j)	5.96	1.048	17.9	(A, B, C), (D, E)
Downs (L4)								
Dhajala (H3)	50 (n)	50 $\pm$ 8.4(b)	45 $\pm$ 9(n)	270 $\pm$ 32(n)	3.17(m)	1.09(m)	8.3(b)	(A, B, C, D, E)
Eitter (H)		4.2 $\pm$ 1.3(h)	57 $\pm$ 2(p)	158 $\pm$ 9(j)	0.77	1.05	1.8	(A), (B, C, D, E)
Eva (L5)		170 $\pm$ 25(h)	42 $\pm$ 2.6(o)	350 $\pm$ 18(o)	17.29(u)	1.09	42	(A, B, C), (D, E)
Edmonson (L6)		33 $\pm$ 8.8(h)	55 $\pm$ 3(o)	225 $\pm$ 23(o)	10.20(q)	1.07	26.7	(A, B, C), (D, E)
Finney (L5)	12 (a)	160 $\pm$ 16(h)	50.2 $\pm$ 3.3(o)	375 $\pm$ 38(o)	7.13	1.15	17.5	(A, B, C), (D, E)
Harrisonville (L)		240 $\pm$ 43(k)	68 $\pm$ 2.4(s)	355 $\pm$ 24(t)	2.58	1.058	6.2	(B, C), (A, D, E)
Keyes (L6)	30 (d)	50(f)	52(f)	393(f)	9.0	1.124	21	(A, B, C, D, E)
		3(f)	68(f)	480(f)	10.35	1.085		(A, B, C, D, E)
Madhipura (L)	6.5 (e)	600(e)	36 $\pm$ 4(h)	280(e)	5.37	1.149(l)	15	(A, B, C, D, E)
Shaw (L)		4.2 $\pm$ 1.4(h)	19.3 $\pm$ 1.4(o)	43 $\pm$ 7(j)	0.29(g)	1.12(g)	0.8	(A, B, C, D, E)
St. Severin (LL)	25 (d, e)	210(f)	65(f)	381(f)	4.8(f)	1.13(f)	11.2	(A, B, C, D, E)
		10(f)	79(f)	420(f)	5.3(f)	1.10(f)		(A, B, C, D, E)
Udaipur (H)	9 (e)	400(e)	38 $\pm$ 6(h)	320(e)	10.85(l)		22	(A, B, C, D, E)

(a) Bhandari *et al* (1980a), (b) Bhandari *et al* (1986), (c) Bagolia *et al* (1978), (d) Bhandari and Poddar (1982), (e) Bhattacharya *et al* (1980), (f) Bhandari (1984), (g) Müller *et al* (1981), (h) this work, (i) Nishizumi *et al* (1980), (j) Englert and Herr (1978), (k) Bhandari *et al* (1980b), (l) Rao M. N. (unpublished), (m) Gopalan *et al* (1977), (n) Bhandari *et al* (1978), (o) Heimann *et al* (1974), (p) Herpers and Englert (1983), (q) P. Englert (private communication), (s) Herpers, *et al* (1968), (t) Herpers and Englert (1983), (u) For data and references see Schultz and Kruse (1983), (v) Measurements [A (track density), B (^{26}Al), C (^{53}Mn), D (^{21}Ne), E ($^{22}\text{Ne}/^{21}\text{Ne}$)] made in the same samples or in samples located at approximately the same shielding depth are grouped together within parentheses.

having multiple exposure and using the criteria developed earlier (Bhandari and Potdar 1982), we confine the study to chondrites with single, one-stage exposure history. A three-isotope (^{26}Al - ^{53}Mn - ^{21}Ne) diagram is proposed which is sensitive to cosmic ray intensity variation. The position of data points in such a diagram allows us to determine the time variation in cosmic ray flux during the past 10^7 years.

2. Experimental details and data sets

We have selected a few meteorites which have already been extensively studied. These are listed in table 1. Track densities in many of these meteorites was not known. This was measured in minerals, olivines and pyroxenes, following standard procedures described earlier (Bhandari *et al* 1972). The cosmic ray track density in olivine is always smaller than in pyroxenes by a factor of two to three. A value of 2.9 has been used here to obtain equivalent density in pyroxenes for the purpose of determining shielding depths from the calculated track production profiles. Samples from Bansur, Madhipura, Udaipur and St. Severin were taken from the cores which have been studied earlier by Bhattacharya *et al* (1980). Samples of Keyes chondrite from different cores were kindly provided by Dr D D Bogard. Aliquots of some of the other meteorites studied here were obtained from Dr P Englert. The track density data in all these samples are listed in table 1. In case of St. Severin the track density, ^{53}Mn , ^{21}Ne and $^{22}\text{Ne}/^{21}\text{Ne}$ have been measured along several cores (Tamhane 1972; Lal *et al* 1969; Bhattacharya *et al* 1980; Englert and Herr 1980; Schultz and Signer 1976). Some ^{26}Al measurements have been made (Marti *et al* 1969) but a depth profile is not available. We have therefore measured ^{26}Al in some samples of core CB4 and A III on a low level gamma-gamma coincidence detector assembly using two $7.5\text{ cm} \times 7.5\text{ cm}$ NaI (Tl) activated crystals with a plastic scintillator active anticoincidence shield. The background of the system is 1.6×10^{-3} cpm and β^+ counting efficiency is 2.3% for a typical sample of 40 g. The

Figure 1. Preatmospheric depth profile of ^{26}Al in St. Severin core samples (Effective radius, $R_E = 25\text{ cm}$). The shielding depths are calculated from the observed track density. Sample number is marked on each point.

samples were powdered, filled in identical boxes and counted. The ^{26}Al profile is shown in figure 1. The $^{26}\text{Al}/^{53}\text{Mn}$ ratio in five samples lie in the range 0.16 ± 0.02 . In case of Keyes, measurements of all these isotopes are available (Cressy 1975; Englert and Herr 1978; Wright *et al* 1973). The $^{26}\text{Al}/^{53}\text{Mn}$ ratio in four samples of Keyes chondrite is 0.13 ± 0.01 . We use the ^{26}Al , ^{53}Mn , ^{21}Ne and $^{22}/^{21}\text{Ne}$ data in Dhajala samples obtained in our laboratory (Bhandari *et al* 1978; Gopalan *et al* 1977; Potdar 1981; Bagolia *et al* 1978). The $^{26}\text{Al}/^{53}\text{Mn}$ ratio in one sample measured is 0.16. In case of Keyes and St. Severin, the best fit curves drawn through the depth profiles summarised earlier (Bhandari 1984) have been used and two data sets, one corresponding to the near surface sample and other near the deepest point analysed are listed in table 1. In other cases we have used data reported by Müller *et al* (1981) and Englert and Herr (1978) as summarised by Moniot *et al* (1983). Shielding depths and approximate preatmospheric

Figure 2. $^{26}\text{Al}/^{21}\text{Ne}$ vs $^{53}\text{Mn}/^{21}\text{Ne}$ correlation diagram for production ratios discussed in the text are plotted for various assumed changes in cosmic ray flux. $\phi_{T_2} = \phi_{T_1}$ represents the correlation curve for constant cosmic ray flux. If the cosmic ray flux has increased during the past 2 million years, the correlation curve will be below and if the flux has decreased, it will lie above the $\phi_{T_2} = \phi_{T_1}$ curve. B: Barwell, BA: Bansur, BR: Bruderheim, D: Dhajala, E: Etter, EV: Eva, ED: Edmonson, F: Finney, K: Keyes, M: Madhipura, N: Nogata, S: St. Severin, SH: Shaw, DD: Dalgety Downs, H: Harrisonville.

sizes, where possible, were calculated using the procedure given in Bhandari *et al* (1980a). All meteorites with long exposure ages (> 3 million years) have been included in table 1 and figure 2. The data on short exposure age meteorites cluster close to the line of constant cosmic ray flux as shown in figure 2 and hence do not indicate any variation in flux. Therefore, only a few typical cases are included.

3. Discussion

We will discuss these data in terms of variation of cosmic ray intensity. However, it may be useful to determine the changes expected in isotope production rates in meteorites for constant and variable fluxes first.

3.1 Isotope production in meteorites

The isotope production rate P_i in a meteorite depends on target element abundance (N_j), excitation function (σ_{ij}) and flux $\phi(E)$ of the primary and secondary nuclear-active particles above the threshold energy E_t . Since $\phi(E)$ depends on the development of nuclear cascade, it changes with the size of the meteorite (R_E , effective radius, if a spheroidal approximation is made) and shielding depth of the sample (X). Further, if the primary cosmic ray flux has been varying with time, time variation of $\phi(E)$ will lead to time variation in P_i . The average production over a time t is then given by

$$P_i = \left[\sum_j \int_t \int_{E_t} N_j \cdot \sigma_{ij}(E) \cdot \phi(E, R_E, X, t) dE dt \right] / \left[\int_0^t dt \right] \quad (1)$$

In case of stable isotopes, e.g. ^{21}Ne , the observed concentration, S_i , is given by

$$S_i = P_i \cdot t, \quad (2)$$

provided that no diffusive losses have occurred from the meteorite. On the other hand, in case of radioactive isotopes, e.g. ^{26}Al or ^{53}Mn , the observed activity, A_k , grows with time,

$$A_k = P_k [1 - e^{-t/T_k}], \quad (3)$$

where T_k is the mean life of the radioisotope. When $t \gg T_k$, i.e. at secular equilibrium $A_k = P_k$ and the exposure age of the meteorite can be calculated by measuring S_i and A_k if the production ratio P_i/P_k is independently known.

$$t = S_i / [A_k [1 - e^{-t/T_k}]^{-1} \times (P_i/P_k)] \quad (4)$$

In chondrites P_i and P_k change significantly (by upto a factor of 2) from sample to sample because of different shielding depths within a meteorite and its size. However, the ratio P_i/P_k is nearly the same (within $\pm 20\%$), if nuclides produced in similar reactions (or with similar nucleon energy) are considered. The exposure ages calculated from any isotope pair e.g. $^{21}\text{Ne}/^{26}\text{Al}$ or $^{21}\text{Ne}/^{53}\text{Mn}$ using (4) should therefore be identical for a given cosmic ray flux.

To determine P_i/P_k we consider three meteorites, St. Severin (LL), Keyes (L) and Dhajala (H), in which ^{53}Mn , ^{26}Al , ^{21}Ne , $^{22}\text{Ne}/^{21}\text{Ne}$ and tracks have been measured (Schultz and Signer 1976; Wright *et al* 1973; Englert and Herr 1980; Cressy 1975; Lal *et*

al 1969; Bhandari *et al* 1978; Bhattacharya *et al* 1980; Potdar 1981). These data form a complete set from which depth profiles have been constructed. The shielding depth in space for each sample was calculated from the track density values using the calculated production rates of Bhattacharya *et al* (1973). These profiles show that the *relative* production rates of ^{26}Al , ^{53}Mn , ^{21}Ne at any depth are nearly constant. The production rates of ^{26}Al : ^{53}Mn : ^{21}Ne are 60 atoms/min kg meteorite: 406 atoms/min kg Fe: 0.41×10^{-8} cc STP/g meteorite. M.Y. and their ratios are not very sensitive to depth or size of the chondrite. These rates are also similar to the model calculations of Bhandari (1981) which give production ratios of 60:415:0.39 (expressed in the units mentioned above) whereas the data of Englert and Herr (1978) in saturated meteorites yield values of 60:392:0.32. Using an altogether different set of experimental data Nishiizumi *et al* (1980) gave the production rates ^{26}Al : ^{21}Ne as 56 atoms/min kg meteorite (for H) or 60 atoms/min kg meteorite (for L and LL): 0.507 ± 0.039 (10^{-8} cc STP/g M.Y.) and ^{53}Mn : ^{21}Ne production rate as 414 atoms/min kg Fe: 0.302 ± 0.013 (10^{-8} cc STP/M.Y.). Moniot *et al* (1983) also favour a low production rate of ^{21}Ne of 0.28 ± 0.02 cc STP/g M.Y. based on ^{10}Be data. Thus there is considerable uncertainty in the production rate of ^{21}Ne . Here we adopt the production ratios in the proportion 60:415:0.39 given by Bhandari (1981) which depend slightly on target composition, shape and size of the meteoroid and shielding depth, though these effects are negligible, < 20% if we confine to ordinary chondrites. In case the cosmic ray flux changes without a significant change in the shape of the energy spectrum, the production ratio is given by

$$P_i/P_k = R_{ik} \bar{\phi}_i/\bar{\phi}_k, \quad (5)$$

where R_{ik} is the production ratio of the isotope for some arbitrary flux and $\bar{\phi}$ represents average flux for the period over which the isotope production is effective.

$$\text{Therefore } t = S_i/[A_k \{1 - e^{-t/T_i}\}^{-1} \cdot R_{ik} \cdot (\bar{\phi}_i/\bar{\phi}_k)] \quad (6)$$

We consider here two values of the exposure age t , calculated from $^{21}\text{Ne}/^{26}\text{Al}$ and $^{21}\text{Ne}/^{53}\text{Mn}$. In case the average cosmic ray flux has been the same over the mean life of ^{26}Al and ^{53}Mn , the two values of t should be the same and they should correlate as shown by the curve $\phi_{T_1} = \phi_{T_2}$ in figure 2. This figure also shows the correlation of t for some arbitrary change in flux in the past over periods (in million years) indicated by the suffix to ϕ . If the flux (ϕ_2) increases by a factor of two over the past two million years, compared to the flux (ϕ_{T_1}) before, i.e. $\phi_2 = 2\phi_{T_1}$, the exposure age correlation curve lies below the $\phi_{T_1} = \phi_{T_2}$ line whereas if the flux reduces to half, i.e. $\phi_2 = 0.5 \phi_{T_1}$, the curve lies above the $\phi_{T_1} = \phi_{T_2}$ line. The curves are significantly different for a factor of 2 or 3 variation in flux over 1 or 2 M.Y. period as shown in this figure so that such variations can be easily detected by comparing experimental data with these curves.

3.2 Complex exposure history of meteorites

It is assumed in the foregoing discussion that the meteorites are exposed to cosmic rays during their entire exposure period in a fixed geometry i.e. X and R_E do not alter with time. As far as the isotope production is concerned, change of X or R_E is equivalent to change of effective nucleon flux. Inherent in the mechanism by which meteorites are produced, such as in collisions of different bodies of the solar system (asteroids, for example), it is quite likely that the object is first irradiated to cosmic rays

buried in the asteroidal regolith (stage T_1) and then exposed as an independent body in space as a meteoroid (stage T_2). In such a case the nucleon flux experienced by the body in stage T_1 would be less than that in stage T_2 . Analysis of the rare gas data has revealed that about 28% of meteorites show comparable exposure in stage T_1 and T_2 (Bhandari 1981). Therefore, in order to deduce information on cosmic ray variations from meteorite data, meteorites with complex exposure history must be excluded.

To distinguish the cases of multiple exposure and genuine cosmic ray variation we use the criterion suggested earlier by Bhandari and Potdar (1982). The tracks have a steep source function with path length of a few millimeters near the surface, increasing with shielding depth to a few centimeters at large depths (Bhattacharya *et al* 1973). The track production is therefore small in stage T_1 and their records in meteorites are mainly due to irradiation in stage T_2 . On the other hand, the spallogenic isotope production increases with depth, initially as a consequence of production of secondary nucleons, and then decreases with a path length of about 50 cm in chondrites (Arnold 1961; Bhandari and Potdar 1981; Reedy and Arnold 1972). The isotope production, for example of ^{21}Ne , therefore would be considerable in both the stages T_1 and T_2 . Furthermore in case of spallation neon isotope ratio $^{22}\text{Ne}/^{21}\text{Ne}$, it has been found that the ratio is sensitive to shielding depth (Bhandari and Potdar 1982; Eberhardt *et al* 1966; Cressy and Bogard 1976). In H chondrites, the ratio is about 1.24 near the surface

Figure 3. Track production rate vs $[^{22}\text{Ne}/^{21}\text{Ne}]_{\text{sp}}$ correlation diagram for various effective radii R_E , following Bhandari and Potdar (1982). The experimental data plotted here are given in table 1. Symbols are the same as in figure 2.

going down to about 1 at a depth of 2 meters. Track production rate vs $[^{22}\text{Ne}/^{21}\text{Ne}]_{\text{sp}}$ correlation diagram for single-stage exposure has been constructed for meteoroids of various sizes (figure 3). Any departure from the expected correlation for a particular size of the meteoroids can be taken to indicate multiple exposure history (Bhandari and Potdar 1982). We use this criterion in the following analysis to identify meteorites with single and multiple exposures. The track density and $[^{22}\text{Ne}/^{21}\text{Ne}]_{\text{sp}}$ for all the meteorites are plotted in figure 3. It is clear from this figure that there are six meteorites, namely, Bruderheim, Finney, Dhajala, Keyes, Shaw and St. Severin in which all the data are consistent with single stage exposure. Other meteorites such as Edmonson, Etter, Dalgety Downs and Harrisonville do not satisfy the criteria of single-stage exposure whereas Barwell and Eva are probably consistent with single exposure within the errors of measurement. The first stage exposure, T_1 , in these two cases, if any, has therefore been negligibly small.

3.3 Analysis of data

If we consider all the data sets, we find that the points lie on various curves, some on the line of constant flux $\phi_{T_1} = \phi_{T_2}$ and others like Dalgety Downs, Edmonson and Nogata towards the line representing $\phi_2 = 2\phi_{T_1}$ (figure 2). Clearly if there is a genuine cosmic ray flux variation, all meteorites must show this effect and the data points must lie on the line characteristic of the change of flux. This is not the case. All the well studied meteorites with single-stage exposures e.g. Bruderheim, Dhajala, Keyes and St. Severin lie close to the curve of constant cosmic ray flux $\phi_{T_1} = \phi_{T_2}$. The meteorites which deviate from this curve are those having complex exposure as indicated by figure 3. Therefore one can conclude that the apparent departure of data from constant flux line in figure 2 is only due to complex exposure of some of the meteorites and not due to increase in the galactic cosmic ray intensity 2 million years ago. Thus the isotope data in meteorites having simple exposure history do not indicate any variation over the past 1 or 2 million years compared to the intensity over the past 10^7 years.

References

- Arnold J R, Honda M and Lal D 1961 *J. Geophys. Res.* **66** 3519
 Bagolia C, Doshi N, Lal D and Sears D W 1978 *Nucl. Track Detection* **2** 29
 Bhandari N 1981 *Proc. Indian Acad. Sci. (Earth Planet. Sci.)* **90** 359
 Bhandari N 1984 *Lunar and planetary Science* (Houston: Lunar Planet Sci. Inst.) vol. 15, p. 50
 Bhandari N, Goswami J N, Lal D, MacDougall D and Tamhane A S 1972 *Proc. Indian Acad. Sci.* **A76** 27
 Bhandari N, Bhattacharya S K and Somayajulu B L K 1978 *Earth Planet. Sci. Lett.* **40** 194
 Bhandari N and Potdar M B 1982 *Earth Planet. Sci. Lett.* **58** 116
 Bhandari N, Lal D, Rajan R S, Arnold J R, Marti K and Moore C B 1980a *Nucl. Tracks* **4** 213
 Bhandari N, Lal D, Nautiyal C M, Padia J T, Potdar M B, Rao M N and Venkatesan T R 1980b *Meteoritics* **15** 265
 Bhandari N, Potdar M B and Suthar K M 1986 *Proc. Indian Acad. Sci. (Earth Planet. Sci.)* **95**
 Bhattacharya S K, Goswami J N, Gupta S K and Lal D 1973 *Moon* **8** 253
 Bhattacharya S K, Imamura M, Sinha N and Bhandari N 1980 *Earth Planet. Sci. Lett.* **51** 45
 Cressy P J 1975 *J. Geophys. Res.* **80** 1551
 Cressy P J and Bogard D D 1976 *Geochim. Cosmochim. Acta* **40** 749
 Eberhardt P, Eugster O, Geiss J and Marti K 1966 *Z. Naturforsch.* **21** 414

- Englert P and Herr W 1978 *Geochim. Cosmochim. Acta* **42** 1635
- Englert P and Herr W 1980 *Meteoritics* **15** 288
- Evans J C and Rancitelli L A 1980 in *Catalog of antarctic meteorites 1977–78* (eds) U B Marvin and B Mason (Washington: Smithsonian Institution) p. 45
- Fleischer R L, Price P B, Walker R M, Maurette M and Morgan G 1967 *J. Geophys. Res.* **72** 355
- Gopalan K, Rao M N, Suthar K M and Venkatesan T R 1977 *Earth Planet Sci Lett* **36** 341
- Heimann M, Parekh P P and Herr W 1974 *Geochim. Cosmochim. Acta* **38** 217
- Herpers U, Herr W and Wölfe R 1968 in *Meteorite research* (Dordrecht: D Reidel) p. 387
- Herpers U and Englert P 1983 *J. Geophys. Res.* **88** Suppl. B316
- Lal D, Lorin J C, Pellas P, Rajan R S and Tamhane A S 1969 in *Meteorite research* (ed.) P M Millman (Dordrecht: D Reidel) p. 275
- Marti K, Shedlovsky J P, Lindstrom R M, Arnold J R and Bhandari N 1969 in *Meteorite research* (ed.) P M Millman (Dordrecht: D Reidel) p. 246
- Moniot R K, Kruse T H, Tuniz C, Savin W, Hall G S, Milazzot T, Pal D and Herzog G H 1983 *Geochim. Cosmochim. Acta* **47** 1887
- Müller O, Hampel W, Kirsten T and Herzog G F 1981 *Geochim. Cosmochim. Acta* **45** 447
- Nishiizumi K, Regnier S and Marti K 1980 *Earth Planet. Sci. Conf.* **50** 156
- Potdar M B 1981 *Nuclear interactions of the solar and galactic cosmic rays with interplanetary materials* Ph.D. thesis, University of Gujarat
- Reedy R C and Arnold J R 1972 *J. Geophys. Res.* **77** 537
- Schultz L and Kruse H 1983 *Helium, neon and argon in meteorites: A data compilation* (Mainz: Max Planck Institut für Chemie)
- Schultz L and Signer P 1976 *Earth Planet. Sci. Lett.* **30** 191
- Shima M, Murayama S, Okada A, Yabuki H and Takaoka N 1983 *Meteoritics* **18** 87
- Wright R J, Simms L A, Reynolds M A and Bogard D D 1973 *J. Geophys. Res.* **78** 1308