

U-Pb age of the Diana Complex and Adirondack granulite petrogenesis

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U-Pb isotopic analyses of eight single and multi-grain zircon fractions separated from a syenite of the Diana Complex of the Adirondack Mountains do not define a single linear array, but a scatter along a chord that intersects the Concordia curve at 1145 ± 29 and 285 ± 204 Ma. For the most concordant analyses, the $^{207}\text{Pb}/^{206}\text{Pb}$ ages range between 1115 and 1150 Ma. Detailed petrographic studies revealed that most grains contained at least two phases of zircon growth, either primary magmatic cores enclosed by variable thickness of metamorphic overgrowths or magmatic portions enclosing presumably older xenocrystic zircon cores. The magmatic portions are characterized by typical dipyrarnidal prismatic zoning and numerous black inclusions that make them quite distinct from adjacent overgrowths or cores when observed in polarizing light microscopy and in back-scattered electron micrographs. Careful handpicking and analysis of the “best” magmatic grains, devoid of visible overgrowth of core material, produced two nearly concordant points that along with two of the multi-grain analyses yielded an upper-intercept age of 1118 ± 2.8 Ma and a lower-intercept age of 251 ± 13 Ma. The older age is interpreted as the crystallization age of the syenite and the younger one is consistent with late stage uplift of the Appalachian region. The 1118 Ma age for the Diana Complex, some 35 Ma younger than previously believed, is now approximately synchronous with the main Adirondack anorthosite intrusion, implying a cogenetic relationship among the various meta-igneous rocks of the Adirondacks. The retention of a high-temperature contact metamorphic aureole around Diana convincingly places the timing of Adirondack regional metamorphism as early as 1118 Ma. This result also implies that the sources of anomalous high-temperature during granulite metamorphism are the syn-metamorphic intrusions, such as the Diana Complex.

1. Introduction

As granulites are generally believed to constitute a major portion of the lower continental crust, a better understanding of granulite petrogenesis can be expected to aid in improved petrologic and geochemical models for the evolution of continental crust. Geothermo-barometric studies of granulite terrains suggest temperatures in the range of 700° to 900°C and pressures of 5 to 10 kbar, implying anomalously high temperatures and deep burial during granulite facies metamorphism. These

results clearly indicate some considerable thermal perturbations in the deep crust that are different and locally steeper than the normal ambient continental geothermal gradient (e.g. Ganguly *et al* 1995).

The Adirondack Mountains of New York have been an important area for the study of massif-type anorthosites and associated granulites for formulating and testing many of the ideas of petrologic, thermal-barometric, geochronological, isotopic, and geochemical aspects of granulite facies metamorphism. It is generally believed

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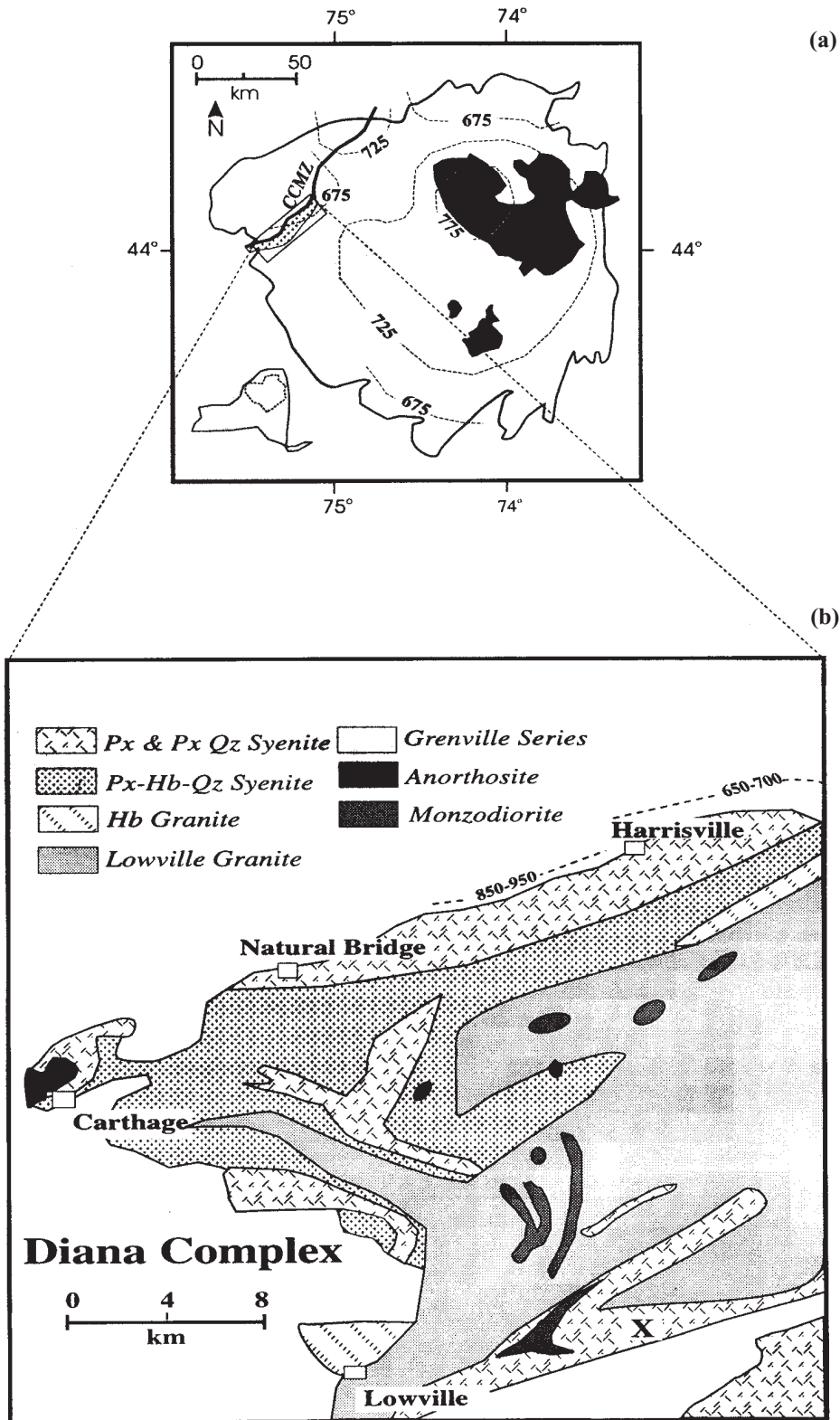


Figure 1. (a). Outline of the Adirondack Mountains, New York, showing the anorthosite massifs (in black) in the central highlands and the Carthage-Colton Mylonite Zone (CCMZ) to the northwest. The CCMZ demarcates the upper amphibolite-facies rocks to the northwest from the higher granulite-grade rocks to the southeast. The location of the Diana Complex is shown in the rectangular box. The contours represent the temperatures of regional metamorphism in degree Celsius, determined from mineralogic thermometers as summarized by Valley *et al* (1994). (b). Simplified lithologic map of the Diana Complex, after Buddington (1939), showing the location of the hornblende-pyroxene syenite (cross) from which the different zircon fractions were separated and analyzed in the present study for U-Pb age determination. The broken lines above Harrisville represent isotherms in degree celsius for contact metamorphism of the adjacent metasediments, as discovered by Powers and Bohlen (1985).

that the north-northeast trending structure within the Adirondack dome (figure 1a) exposes a section of the middle Proterozoic crust composed of upper amphibolite to granulite facies rocks. Granitic, charnockitic and anorthositic gneisses of the Adirondack highlands along with some intervening metasediments have also been considered as a plausible section of the deep continental crust. High-grade regional metamorphism associated with the Grenville event (ca 1100 Ma) has pervasively affected the entire Adirondack region, and the progressive transition from upper amphibolite to granulite-facies metamorphism across the major paragneisses surrounding the anorthosite and other orthogneisses has been well-documented in the northwest Adirondacks (e.g., Edwards and Essene 1988). Extensive geothermo-barometric studies of Adirondack metamorphism (e.g., Bohlen *et al* 1985; Newton 1985) have documented strong correspondence between the highest temperatures and pressures and the outline of the main anorthosite massif in the Adirondack highlands (figure 1a). These results are summarized by several workers (e.g. Newton 1985; Bohlen *et al* 1985) based on the investigations of Buddington (1963); Bohlen and Essene (1977); Johnson and Essene (1982); Newton (1983) and others. These pressure-temperature estimates, which form a "bull's-eye" pattern centered around the anorthosite massif (figure 1a), may have important implications concerning the timing of anorthosite intrusion, metamorphism, and the uplift of the Adirondack highlands.

During the past four decades, geochronological evidence for Adirondack metamorphism and anorthosite intrusion has been interpreted with broadly different results. The first definitive U-Pb isotopic study of zircons by Silver (1969) indicated no evidence of age of intrusion significantly older than 1115 Ma among the orthogneisses of the Adirondack highlands. In addition, Silver's data demonstrated an overall duration of 80 ± 25 Ma for the post-anorthosite episode of metamorphism. An internal mineral and whole-rock Sm-Nd isotopic study for garnetiferous anorthosites of the Snowy Mountain massif in the central Adirondacks gave the time of formation of the garnets and possibly of the minimum age of crystallization of the anorthosites at 1100 Ma (Basu and Pettingill 1983), consistent with Silver's interpretation of the U-Pb data in zircons. Although these studies, among others, indicated the synchronicity of anorthosite emplacement, and the continuity and duration of metamorphism following plutonism, many Adirondack workers during the last two decades (e.g., Ashwal and Wooden 1983, Valley 1985; McLelland and Isachsen 1986; and others) accepted a broad window of 400 Ma between the timing of magmatic and metamorphic events

in the Adirondacks. This view was primarily based on the Sm-Nd systematics results of Ashwal and Wooden (1983) that yielded isochron ages between 1288 and 919 Ma for anorthosite emplacement and the following metamorphic episodes. In the light of new U-Pb data (Chiarenzelli *et al* 1987; McLelland *et al* 1988; McLelland and Chiarenzelli 1989, 1990; Mezger *et al* 1991) of meta-igneous and meta-sedimentary rocks of the Adirondack region, Silver's (1969) broad interpretations have now been vindicated and it appears likely that Ashwal and Wooden's (1983) results reflected disturbed Sm-Nd systematics (McLelland *et al* 1988). However, controversies still exist on whether the rocks adjacent to the anorthosites experienced polymetamorphism (e.g., Valley *et al* 1990) or, that the central anorthosite massifs intruded contemporaneously with the anorthosite-mangerite-charnockite-granite (AMCG) suite (Emslie and Hunt 1989; McLelland and Chiarenzelli 1990; Chiarenzelli and McLelland 1991), followed by granulite-grade metamorphism.

In this paper, we have attempted high-precision U-Pb dating of single and multiple grains of zircons from the Diana Syenite Complex of the northwest Adirondacks. Our purpose is to address some of the above questions of precise age relationships between timing of intrusion and metamorphism. We provide detailed petrographic information on the zircons analyzed (table 1) and, in conjunction with the U-Pb isotopic analyses of the zircons, we demonstrate that many of the zircon fractions from the same hand-specimen of a syenite contain either primary magmatic grains with metamorphic overgrowths, or older xenocrystic core material rimmed by magmatic portions, as well as some rare purely "magmatic" grains. It is the analyses of the rare magmatic grains that produced nearly concordant ages, consistent with the expected field, petrologic, geothermo-barometric, and other geochronological studies of Adirondack granulite lithology.

2. Diana Syenite Gneiss Complex of northwest Adirondacks

The Diana Quartz Syenite Gneiss Complex is composed of a variably metamorphosed rock series ranging from augite-hypersthene syenite through augite syenite, augite-hornblende syenite, hornblende-quartz syenite, granosyenite and granite (Buddington 1939, 1963). Buddington (1939) and Hargraves (1969) interpreted the Diana Complex as a metamorphosed, differentiated quartz-syenite igneous intrusive exposed along the Carthage-Colton Mylonite Zone (CCMZ), speculating that the CCMZ represented an overthrust fault zone marked by mylonitization with the

Table 1. *Petrographic descriptions of zircon fractions separated from the Diana Complex syenite.*

Sample	Wgt. [‡] (mg)	Form	Color	Features	Inheritance?
1 (single grain)	0.040	Euhedral yellowish-pink	Clear, slight very cracked	Not whole xtal	No cores observed
2 (single grain)	0.029	Euhedral dipyramidal	Clear ?	Whole xtal, very cracked	Some inclusions?
3 (single grain)	0.025	Euhedral- subhedral	Clear ? some reddish-brown stains	Whole xtal, some cracks	<u>Possible core observed.</u>
4 (six grains)	0.098	Euhed-subhed, elongate	Clear ? some yellowish-red stains	Mostly whole xtals, some cracks	No cores observed
5 (single grain)	0.012	Euhed-subhed	Clear ? no discolorations	Mostly whole xtals, one small black inclusion	No cores observed
6 (single grain)	0.017	Euhed-subhed	Mostly clear, no discolorations	Whole xtal, cracks, small black inclusions	No cores observed
7 (single grain)	0.014	Subhedral	Clear with some red stains (Fe ?)	Whole xtal, some cracks	No cores observed
8 (nine grains)	0.169	Euhed-subhed	Some discolored	Some whole – some not; many cracks and small black inclusions	<u>Two with possible cores</u>
9 (single grain)	0.038	Euhed-subhed, 1/w = 5–6, elongate	Clear, no discolorations	Whole xtal, some cracks, small black inclusions	No cores observed
10 (single grain)	0.023	Euhed-subhed, 1/w = 6–7, elongate	Clear, no discolorations	Whole xtal, fewer cracks, small black inclusions	No cores observed, but <u>one red inclusion?</u>
11 (six grains)	0.084	Euhed-subhed, most elongate, two shorter	Clear, no discolorations	3 whole – 3 partial xtals, cracks, small black inclusions	No cores observed,
12 (eight grains)	0.144	Euhed-subhed, stubby, 1/w = 1–3	Turbid, more discolored than others	5 whole – 3 partial xtals, many cracks and inclusions	<u>Cores likely, one with a distinct frosty core</u>

Grenville belt in the northwest thrust over the main Adirondack massif to the southeast. Figure 1(a) shows the location of the CCMZ in the northwest Adirondacks, including the extent of the Diana Complex within it to the southwest, and its relationship with the major anorthosite domes in the central highlands. The concentric patterns of isothermal contours in figure 1(a) are those of metamorphic temperatures estimated from various mineralogical thermometers as summarized by Bohlen *et al* (1985). The CCMZ is an important petro-tectonic marker in the sense that it separates the upper amphibolite facies rocks to the northwest from the granulite-facies rocks in the eastern highlands.

Figure 1(b) is a simplified geological map of the Diana Complex after Buddington (1939), who considered the various lithologies as belonging to a closely folded, isoclinally overturned, gravity-stratified plutonic sheet of rocks with an underlying chilled-facies of pyroxene syenite. The Diana syenites are quite similar in mineralogy to the hypersthene granulites of the Adirondack highlands. The Diana pyroxene syenites may belong to the charnockitic or mangeritic rock series surrounding the main anorthosite massifs in the highlands as suggested by DeWaard (1969), although the average Diana syenite contains only 0.7 per cent hypersthene in the mode (Buddington 1939). It is interesting that Buddington's (1939) petrologic description of the various lithologic units of the Diana Complex (figure 1b) can be matched precisely by the volumetrically dominant, bimodal suite of meta-igneous rocks of the Adirondack highlands, including anorthosite, mangerite, charnockite, and granite (the AMCG suite).

A precise information on the primary crystallization age of the Diana Igneous Complex would be highly desirable for at least two reasons. First, the Diana Complex is lithologically similar and possibly genetically related to other mangeritic and charnockitic complexes in direct contact with the anorthosites of the central Adirondacks (Buddington 1939; DeWaard 1969; Wiener 1983). Thus, the age of primary crystallization of the Diana Syenites would correlate with that of the granulitic meta-igneous rocks in the central Adirondacks. Second, the discovery by Powers and Bohlen (1985) of a contact metamorphic aureole, decreasing from 850–950°C along the Diana-metasediment contact to 650–700°C within 2–3 km from the contact (figure 1b), has important implications for Adirondack metamorphism. Powers and Bohlen (1985) interpreted their observation as a "preserved syn-regional metamorphic contact aureole," because it is highly unlikely for the observed contact aureole to have survived the regional granulite-grade metamorphism advancing later from the central Adiron-

dacks. This interpretation is consistent with the increased degree of partial melting of the metasediments towards the Diana syenite contact (Powers and Bohlen 1985). The above considerations prompted us to undertake the current study to determine the age of crystallization of the Diana Complex, and to infer the relative timing of granulite metamorphism in the Adirondacks, and to explore a possible cogenetic relationship among the various meta-igneous lithologies.

3. Analytical methods

Zircons from a sample of syenite within the Diana Complex (marked with a cross in figure 1b) were separated using conventional methods of crushing, sieving, density and magnetic separations. This rock is a foliated medium-grained, equigranular, alkali-hornblende-pyroxene syenite, collected from a road-cut approximately 1 km east of Crystal Dale. Detailed petrographic studies of the entire unbiased zircon population revealed that most grains contained at least two phases of zircon growth, either primary magmatic portions enclosed by variable thicknesses of metamorphic overgrowths or magmatic portions enclosing presumably older xenocrystic zircon cores (figure 2a). The magmatic portions are characterized by typical dipyratidial prismatic zoning and numerous small black inclusions that make them quite distinct from adjacent overgrowths or cores when observed through plane-polarized or crossed nicols (figure 2b) or in back-scattered electron micrographs. A significant percentage of grains exhibited all three components. Careful handpicking of the "best" magmatic euhedral grains, devoid of visible overgrowth or core materials with a larger length to width ratios, was attempted in this study in order to minimize two-component mixing in our analyses. Individual zircon grains, both whole and fragmented, as well as multi-grain fractions were hand-picked from a least-magnetic, heavy concentrate, primarily based on their clarity, form, and lack of inclusions as well as observable overgrowths or cores. A list of petrographic observations made under a reflected-light, binocular microscope prior to dissolution is given in table 1. In general, only grains (whole or fragmented) that were relatively colorless, inclusion-free, core-free with good dipyratidial prismatic form (euhedral to sub-hedral), length to width (l/w) ratios greater than 3 (except fraction 12), and with some indication of magmatic growth zones were selected. Unfortunately, most of the magmatic-looking grains also had numerous small black inclusions that were not always discernible from xenocrystic cores. Several of the fractions contained possible xenocrystic

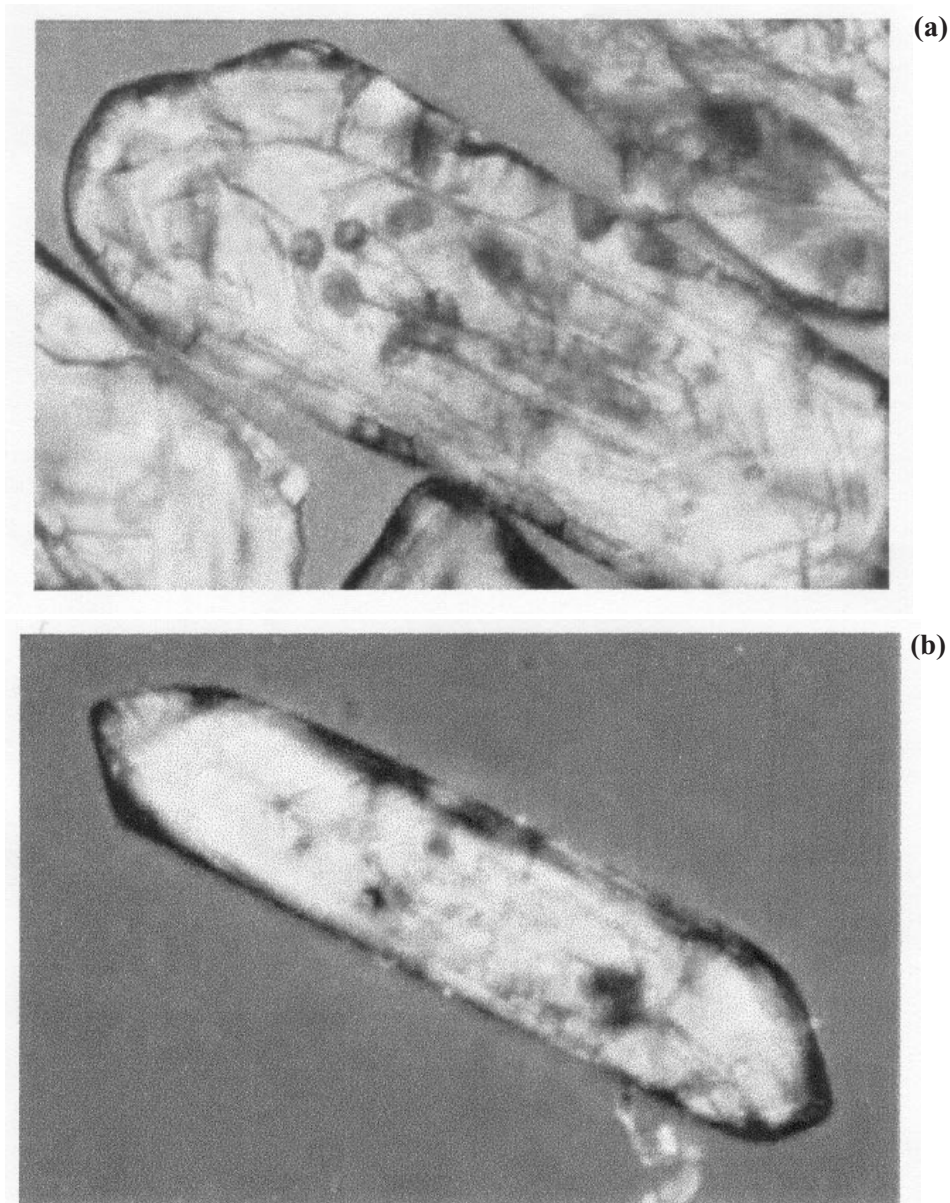


Figure 2. (a). Polarized light photomicrograph in crossed nicols of a zircon population from the syenite, showing multiple overgrowths, heterogeneous cores, and numerous other mineral inclusions. Analyses of these zircons show them to be widely discordant with variable amounts of inherited Pb components in the U-Pb isotope systematics diagram (figure 3). Polarizing light microscopic observations of these zircons revealed the numerous inclusions as apatite and zircons. Notice also the stubby habits of the zircons. Long dimension of photomicrograph is 0.75 mm. (b). Polarized light micrograph of a single, magmatically zoned, euhedral zircon in crossed nicols from the same syenite in which the more common zircons look like those in figure 2(a). Notice smooth, long, slender, euhedral crystal (part of the grain is under the epoxy and not visible in this view) without inclusions and metamorphic cores. We consider this type of grain as “magmatic” and they give near-concordant U-Pb ages. Our conclusions of the primary crystallization age of the Diana Complex is based on the analyses of several grains of zircons similar to this crystal. Long dimension of photograph is 0.40 mm.

tic core material, including fractions 3, 8, and 12. Also, nearly every grain was highly fractured. Whether this attribute was caused during sample preparation or by some natural process is uncertain.

The analytical methods for the dissolution of the zircons and the extraction of U and Pb are the same as those given in Premo and others (1990), which also includes the multi-sample, vapor-dissolution technique of Krogh (1978). Prior to dissolution, the zircon grains were weighed into PFA Teflon

vials, cleaned in distilled 7N HNO₃, and spiked with a ²⁰⁵Pb-²³³U-²³⁶U dilute tracer solution. Dissolution was achieved in distilled concentrated HF + HNO₃ in a large (6.5-cm-diameter) Parr-type TFE Teflon dissolution vessel in 210°C for about 5 days. U and Pb from individual grains were extracted using anion exchange resins, Pb in an HBr medium and U in a HNO₃ medium. Total analytical blanks (laboratory contamination) measured 25 to 70 pg for Pb and 1 to 3 pg for U (see table 2).

Pb was loaded onto a single Re filament using the silica-gel-phosphoric acid method, and isotopic ratios were measured using either an NBS-type tandem (two-stage) mass spectrometer equipped with an ion pulse counter at the end of the second stage or VG Isomass 54E. Pb isotopic ratios were corrected for mass fractionation and laboratory blank Pb (values given in the footnotes of table 2); and for initial common Pb using values of Stacey and Kramers (1975) for the appropriate age of the sample at $^{238}\text{U}/^{204}\text{Pb} = 9.74$. Uncertainties in isotopic ratio measurements are given in table 2, and are reported at the two-sigma level. Ages were calculated using decay constants from Steiger and Jäger (1977). Concordia intercept ages and errors were determined using the algorithms of Ludwig (1980, 1987), which use the regression approach of York (1969); uncertainties on the ages are reported at the 95% confidence level.

4. Results

The U-Pb isotopic data from the twelve analyses, including the 8 single-grain are given in table 2 and shown on a conventional concordia diagram (figure 3). U concentrations range between 340 and 1000 ppm, not unlike other syenitic zircons. There is a rough correlation between U content and concordancy. The higher the U content, the greater is the degree of discordance, although there are exceptions. The analyses were relatively clean, uncorrected $^{206}\text{Pb}/^{204}\text{Pb}$ values ranged between 1500 and 10,200, so that corrections for common Pb were very slight.

The analyses are neither concordant nor collinear on the Concordia diagram (figure 3). The spread in $^{207}\text{Pb}/^{206}\text{Pb}$ ages (table 2) was from 1057 Ma (fraction #4) to 1159 Ma (fraction #10), although for the most concordant analyses, the ages ranged from 1115 to 1150 Ma. These differences are outside the $^{207}\text{Pb}/^{206}\text{Pb}$ age error limits (table 2), indicating that some of the fractions contain at least two components of zircon growth of different ages. At the beginning of our selections of zircon grains, we purposefully picked the “best” grain first. This grain, in our opinion, was fraction # 1, and was the largest, euhedral, and most clear zircon in the entire population. As it turned out, analysis # 1 is also the most concordant, yielding a $^{207}\text{Pb}/^{206}\text{Pb}$ age of 1119 ± 9 Ma (figure 3).

Whereas other single grains were selected that were “next best”, none except analysis # 7, turned out to be so concordant. Analysis # 7 also had a young $^{207}\text{Pb}/^{206}\text{Pb}$ age of 1115 ± 3 Ma. Because of the close proximity and relatively large errors on these two analyses, a chord between them does not yield a useful age. Analyses # 4 and # 12 are both

batch fractions of six and eight grains, respectively, that were not observed to be of the very best quality, and # 12 possibly included a grain containing xenocrystic core material that happens to lie very near this chord. A regression of analyses #1, #7, #4, and #12 yields an upper-intercept age of 1118 ± 2.8 Ma (figure 3) that we presently interpret as the “best estimate” towards the true age of the syenite from the Diana Complex. The remaining analyses, therefore, are interpreted as containing variable amounts of inherited older Pb due to incorporation of xenocrystic zircon cores within the primary magmatic portion of the grain(s). Pb isotopic compositions of older xenocrystic zircon have higher $^{207}\text{Pb}/^{206}\text{Pb}$ values, so that even some small fragment of this older component will cause the analyses to shift slightly to the right of the 1118-Ma chord (figure 3). Some of these fractions were observed to have xenocrystic cores, but for the most part, they were not obviously visible during handpicking. This situation is not too surprising after viewing the small size and similar color of these cores to other normal black inclusions. Our observations, however, identify the cores by their central nature within the magmatic portion of some grain(s). Growth zones tend to emanate from the cores, which is a likely situation since the xenocrystic cores are thought to act as seeds in order to nucleate normal zircon growth within a magma reservoir. Other chords can be constructed through the analyses, but only a few yield lower intercept ages that are reasonably consistent with probable Appalachian uplift ages between ~200 and 400 Ma. Thus our “best estimate” chord with a lower-intercept age of 251 ± 13 Ma (figure 3) may indeed support this regression as a “best estimate.” Individual analytical points are considered mixtures of older inherited Pb and primary magmatic Pb (and possibly younger overgrown Pb as well) which yield a false crystallization age of ~1145 Ma.

5. Discussion

It is our opinion that the “best estimate” for the crystallization age of the syenite from the Diana Complex is 1118 ± 2.8 Ma based on the observations and reasoning given above. By analyzing only the “best” single magmatic grains we were able to decipher this age. This age is in conflict with a U-Pb zircon age of 1155 ± 4 Ma from a syenite reported by Grant *et al* (1986); however, the details of this zircon age results were never published, and reported only in an abstract (the isotopic data were not published). It may have been determined using relatively large zircon populations (mg-size) to measure the radiogenic Pb isotopic compositions. As we pointed out above, many of the zircon

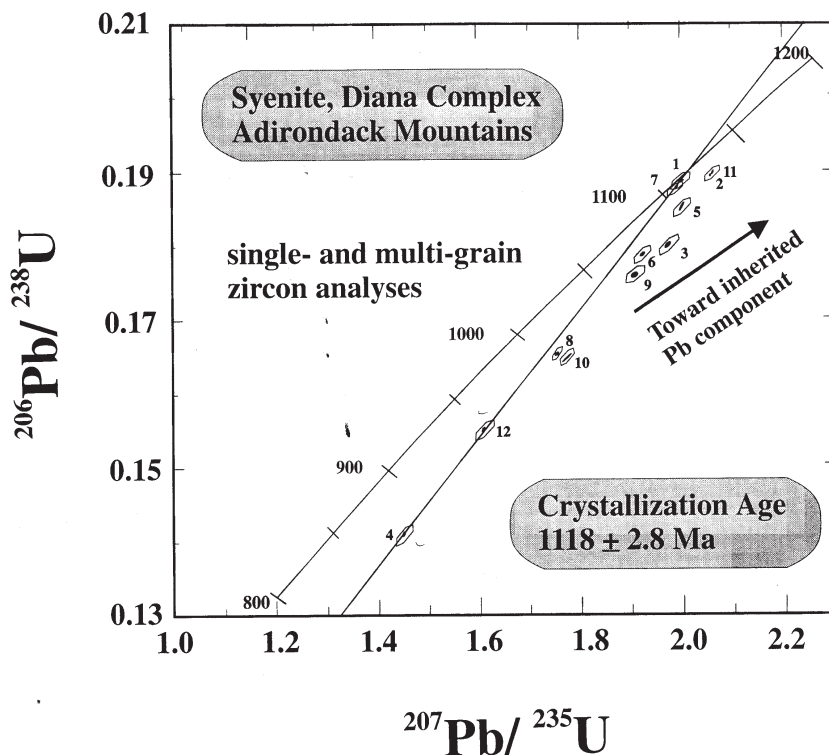


Figure 3. U-Pb concordia diagram plot of the twelve analyses in table 2 of the single and multi-grains of zircons from the syenite of the Diana Complex. The age of 1118 Ma is obtained by the regression of two single-grain (#1 and #7) and two multi-grain (#4 and #12) analyses. See table 1 for petrographic descriptions of the zircons analyzed in this study.

fractions may contain mixtures of two or more zircon components (e.g. overgrowths, magmatic portions, and older xenocrystic cores) that can lead to false (perhaps older) ages. We believe that this is the case here. A reaffirmation of the zircon age from Grant *et al* (1986) was also made by U-Pb mineral age of metamorphic rocks from the contact zone of the Diana Complex (Mezger *et al* 1991), and an age of 1153 ± 3 Ma was reported for garnets in the syn-regional metamorphic contact aureole of the Diana syenite. Our preferred age of crystallization of the Diana Complex at 1118 Ma is some 35 Ma younger than this contact-metamorphic U-Pb garnet age of Mezger *et al* (1991). This discrepancy could be due to the low concentration of U (1.89 ppm) in the garnets and the uncertainty in our knowledge of submicroscopic scale inclusions of older U-Pb bearing minerals in the garnets. The strong similarity between the U-Pb age of zircon by Grant *et al* (1986) and the contact-metamorphic U-Pb age of garnet growth by Mezger *et al* (1991) is indeed puzzling. This is because Grant *et al*'s zircon age is very likely a magmatic crystallization age and, therefore, should be older than the intrusion-induced, contact-metamorphosed garnets, dated by Mezger *et al* (1991). Furthermore, acceptance of the ca 1155 Ma for Diana requires that the U-Pb closure temperatures for zircons and garnets are the same, and that the conductive-convective heat-

ing of the Diana country rocks were almost instantaneous with the intrusion. Both these requirements cannot be accepted because garnet closes at lower temperature with respect to U-Pb diffusion than zircon, and the relatively large 2–3 km distance between the dated garnet and the intrusive contact. Therefore, we prefer our 1118 Ma age for Diana.

It is noteworthy that the 1118 Ma age for the Diana Complex, as indicated by this study, is consistent with Silver's (1969) conclusion regarding the age of intrusion of the anorthosite-charnockite suite, and the more recent assertion by McLelland and Chiarenzelli (1990) on the basis of field evidence and U-Pb dating that the main anorthosite body of the Adirondacks (Marcy Massif) intruded sometime during 1113 to 1138 Ma. Acceptance of the 1118 Ma age of intrusion for the Diana Syenites, as proposed here, independently supports the general conjecture prevalent among Grenville geologists, Adirondacks in particular, of the contemporaneity of the anorthosite-mangerite-charnockite suite of rocks (e.g., McLelland *et al* 1988; Emslie and Hunt 1989). The synchronicity of the Diana Complex with the anorthosite massifs of the Adirondack highlands has important implications for the petrogenesis of the meta-igneous suite, and the following granulite-grade metamorphism. The results of this study and those of the Sm-Nd

systematics of the major meta-igneous rocks of the Adirondacks, including the anorthosites, gabbros, mangerites, charnockites, and granites (Sharma and Basu 2001, in preparation), indicate consanguineous relationship among these rock-types, possibly accompanying assimilation and fractional crystallization processes in a large body of mantle-derived mafic magma emplaced around 1120 Ma in the lower crust. The regional granulite to upper amphibolite grade metamorphism of the Adirondacks was also synchronous with this igneous activity, as indicated by the contact metamorphic aureole around the Diana Complex (Powers and Bohlen 1985) and began as early as 1118 Ma, and continued through 1050 Ma (Basu and Pettingill 1983; McLelland *et al* 1988; Basu *et al* 1989). The regional granulite metamorphism cannot be younger than the age of intrusion of the Diana at 1118 Ma because the younger high-grade regional metamorphism would have completely erased the contact metamorphic aureole adjacent to Diana. The only viable option is to accept the synchronicity of granulite metamorphism with the igneous activity in the Adirondacks because, as we have discussed above, there is no credible evidence for the granulite metamorphism to be older than the intrusion.

Previous zircon age determinations for many meta-igneous rocks from the Adirondack Mountains as well as surrounding orogenic belts have utilized relatively large zircon populations (mg-size) to measure the radiogenic Pb isotopic compositions (Grant *et al* 1986; McLelland *et al* 1988; McLelland and Chiarenzelli 1989; Chiarenzelli and McLelland 1991). We have demonstrated here that it is very possible that many of these zircon fractions may contain mixtures of two or more zircon components (e.g. overgrowths, magmatic portions, and older xenocrystic cores) that can lead to older ages. If the largest, euhedral and most clear grain of the zircon population is also most concordant, reliable age information of the primary crystallization of the meta-igneous rocks can be obtained.

6. Conclusions

We have provided new U-Pb isotopic data in zircons from a syenite of the Diana Complex, northwest Adirondacks. The crystallization age of the Diana Complex, inferred from the "best" magmatic grains of zircons, is nearly concordant at 1118 Ma \pm 2.8 Ma. This age, 35 Ma younger than previously believed, is approximately synchronous with the age of intrusion of the main anorthosite massif in the central highlands, suggesting a cogenetic relationship among the vari-

ous meta-igneous rocks. Because the Diana Complex is partially enclosed by a high temperature contact-metamorphic aureole, the 1118 Ma age is also the time of initiation of Adirondack regional metamorphism which continued through 1100 Ma until 1050 Ma. These geochronological results, consistent with an earlier study by Silver (1969), imply that the anomalously high temperatures for granulite metamorphism in the Adirondack Mountains were caused by the syn-metamorphic intrusions of the anorthosite-mangerite-granite suite of rocks, without multiple periods of heating and cooling, and burial and uplift.

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