

In conclusion, in the present study we have successfully induced adventitious roots from the leaf explants of *A. paniculata*. The adventitious roots were cultured in flask-scale suspension cultures using MS medium supplemented with 2.7 μ M NAA and 30 g/l sucrose. Adventitious root cultures showed higher biomass as well as andrographolide accumulation capabilities. Our study demonstrates the possibilities of production of andrographolides for commercial purposes in a large scale using bioreactor cultures.

1. Sharma, A., Singh, R. T., Sehgal, V. and Handa, S. S., Antihepatotoxic activity of some plants used in herbal formulation. *Fitoterapia*, 1991, **62**, 131–138.
2. Tang, W. and Eisenbrand, G., *Chinese Drugs of Plant Origin*, Springer-Verlag, Berlin, 1992, pp. 97–103.
3. Mishra, P., Pal, N. L., Guru, P. Y., Katiyar, J. C. and Srivastava Tandon, J. S., Antimalarial activity of *Andrographis paniculata* (Kalmegh) against *Plasmodium berghei* NK 65 in *Mastomys natalensis*. *Int. J. Pharmacogn.*, 1992, **30**, 263–274.
4. Saxena, S., Jain, D. C., Bhakuni, R. A. and Sharma, R. P., Chemistry and pharmacology of *Andrographis* species. *Indian Drugs*, 1998, **35**, 458–467.
5. Kumar, R. A., Sridevi, K., Kumar, N. V., Nanduri, S. and Rajagopal, S., Anticancer and immunostimulatory compounds from *Andrographis paniculata*. *J. Ethnopharmacol.*, 2004, **92**, 291–295.
6. Verpoorte, R., van der Heijden, R., ten Hoopen, H. J. G. and Memelink, J., Metabolic engineering of plant secondary metabolite pathways for the production of fine chemicals. *Biotechnol. Lett.*, 1999, **21**, 467–479.
7. Dornenburg, H. and Knorr, D., Strategies for improvement of secondary metabolite production in plant cell cultures. *Enzyme Microb. Technol.*, 1995, **17**, 674–684.
8. Rao, S. R. and Ravishankar, K. A., Plant cell cultures: chemical factories of secondary metabolites. *Biotechnol. Adv.*, 2002, **20**, 101–153.
9. Butcher, D. N. and Connolly, J. D., An investigation of factors which influence the production of abnormal terpenoids by callus cultures of *Andrographis paniculata* Nees. *J. Exp. Bot.*, 1971, **22**, 315–322.
10. Murashige, T. and Skoog, F., A revised medium for rapid growth and bioassays with tobacco tissue cultures. *Physiol. Plant.*, 1962, **15**, 474–497.
11. Jain, D. C., Gupta, M. M. and Saxena, S. K., LC analysis of hepatoprotective diterpenoids from *Andrographis paniculata*. *J. Pharm. Biomed. Anal.*, 2000, **22**, 705–709.
12. Klerk, G. J. D., Arnholdt-Schmitt, B., Lieberei, R. and Neumann, K. H., Regeneration of roots, shoots and embryos: physiological, biochemical and molecular aspects. *Biol. Plant.*, 1997, **39**, 53–66.
13. Gao, X., Zhu, C., Jia, W., Gao, W., Qiu, M., Zhang, Y. and Xiao, P., Induction and characterization of adventitious roots directly from leaf explants of *Panax notoginseng*. *Biotechnol. Lett.*, 2005, **27**, 1771–1775.
14. Wu, C. H., Murthy, H. N., Hahn, E. J. and Paek, K. Y., Large-scale cultivation of adventitious roots of *Echinacea purpurea* in airlift bioreactors for the production of chichoric acid, chlorogenic acid and caftaric acid. *Biotechnol. Lett.*, 2007, **29**, 1179–1182.

ACKNOWLEDGEMENTS. This work was partially funded by the University Grants Commission (Special Assistance Programme), Department of Science and Technology (FIST Programme), Council of Scientific and Industrial Research and National Medicinal Plant Board, New Delhi, India.

Received 24 December 2007; revised accepted 23 December 2008

Climatic influence on radial growth of *Pinus wallichiana* in Ziro Valley, Northeast Himalaya

Santosh K. Shah¹, Amalava Bhattacharyya^{1*} and Vandana Chaudhary²

¹Birbal Sahni Institute of Palaeobotany, 53, University Road, Lucknow 226 007, India

²Department of Science and Technology, New Mehrauli Road, New Delhi 110 016, India

An attempt has been made here to study the climatic influence on variation of tree-ring width (radial growth) of Blue Pine (*Pinus wallichiana* A.B. Jackson) growing in five different sites in and around Ziro Valley, Arunachal Pradesh, Northeast Himalaya. The site chronologies have been evaluated to assess inter-site differences through several statistical analyses, viz. correlation matrices, principal component and hierarchical cluster analysis. Analysis of tree growth–climate relationship suggests that the pre-monsoon precipitation (December–April) is a significant factor influencing the growth of Blue Pine in all these sites.

Keywords: Blue Pine, climatic influence, radial growth, tree ring.

BLUE Pine (*Pinus wallichiana*), a large evergreen conifer tree, is found all along the Himalayas from west Kashmir to east Arunachal Pradesh¹, at altitude generally ranging from 1800 to 3900 m. Typical habitats are mountain screes and glacier forelands and appears as a pioneer species, but it also forms old growth forests as a primary species in mixed forests with deodar (*Cedrus deodara*), spruce (*Picea smithiana*) and fir (*Abies pindrow*) in the temperate belt. In some places beyond 3000 m, it reaches the tree-line and is associated with birch (*Betula utilis*) and juniper (*Juniperus macropoda*)². Blue Pine growing in several sites of the western and central Himalaya has been found suitable for various environmental and ecological applications in tree-ring analyses^{3–9}, but no such analyses are available from the northeastern part of the Himalaya.

In this paper an attempt has been made to analyse tree rings of *P. wallichiana* growing in and around Ziro Valley, Lower Subansiri District, Arunachal Pradesh, Northeast Himalaya, towards understanding its dendroclimatic potential from a new geographical region which is greatly influenced by the southwest monsoon.

Tree-ring samples of Blue Pine were collected from five forest sites characterized by open pine mixed broad leaved forest in and around Ziro Valley. These sites are Doby (DOB), Hari (HAR), Hong (HON), Michi-Rant (MIC) and Raga (RAG; Figure 1). All these sites are

*For correspondence. (e-mail: amalava@yahoo.com)

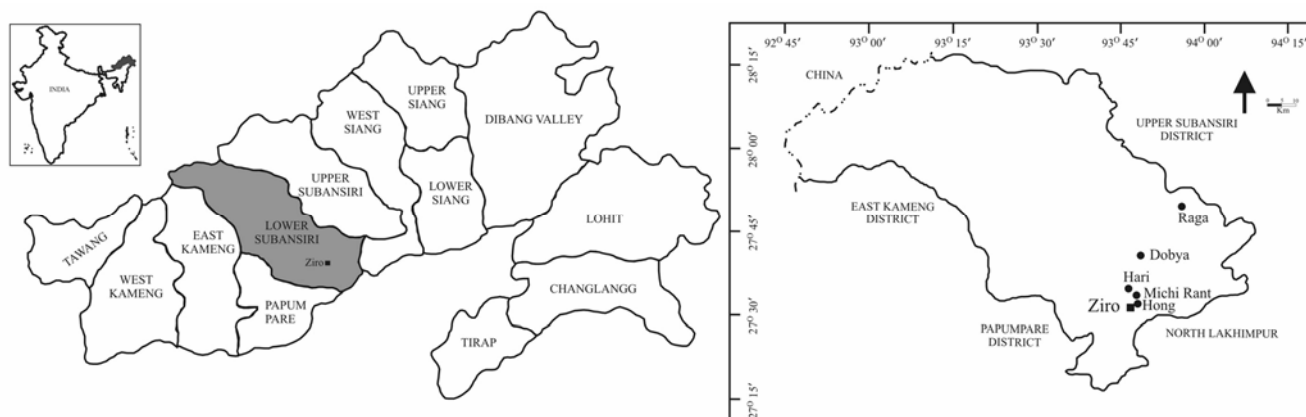


Figure 1. Map showing sampling sites of *Pinus wallichiana* at Ziro Valley, Lower Subansiri District, Arunachal Pradesh.

Table 1. Details of the tree-ring sampling sites from Ziro Valley

Site	Site code	Altitude (m)	Number of trees and cores
Doby	DOB	1320	4 (11)
Hari	HAR	1260	17 (36)
Hong	HON	1245	3 (6)
Michi-Rant	MIC	1255	15 (35)
Raga	RAG	1582	6 (12)

located at latitude and longitude ranging from 27°53'–27°68' and 93°85'–93°89' respectively. These sites are close to human settlements and thus highly disturbed. In HON village trees are confined to steep, rocky slopes, but most of them are found burnt by forest fire. Generally trees are young, except in some sites where several isolated, huge, girthed pine trees are recorded which have been preserved by the villagers in the memory of their ancestors. Otherwise, all old trees have been already logged. These protected trees, mentioned to be about 2000 years old by the local people, were planted by their ancestors after migrating from Tibet, and now these are identified by their names. Hundred cores from 45 trees of Blue Pine have been collected from these five forest sites with the help of Increment borer. Details of the tree-ring samples are given in Table 1.

All the tree cores of Blue Pine were mounted on a wooden frame. The surface of these cores was cut by the sharp razor blade and polished with the help of different grades of sandpaper to enhance the clarity of the cells. Tree rings of these cores were counted under stereo-zoom microscope and were dated to the calendar year of their formation using the typical skeleton plot technique¹⁰. The width of each dated tree-ring in each core was measured using VELMEX measuring system coupled with a computer. These measurements and dates were checked using the computer program COFECHA^{11,12}. Cores that correlated poorly, even after rechecking and corrections (if

there were any errors), were removed from the population prior to chronology development. In order to remove the gross growth trend largely unrelated to climate, and to remove effects of differential mean growth rates among individual trees prior to averaging into the mean chronology, standardization of each series was made. It was done by fitting a growth curve to the ring-width series and then dividing the measured value by the curve value at each year. After applying different detrending options for the standardization to check the improvement of chronology statistics, we have selected cubic spline which is appropriate for trees growing within the interior of the forests and not growing in stressed condition. The detrending for all the series was done using 30-year spline which retains 50% of the variance contained in the measurement series at a wavelength of 30 years. Two versions of chronologies, 'standard' and 'residual', were prepared for each site intended to contain a maximum common signal and minimum amount of noise. In the 'standard' chronology, detrending of measurement series was done first by fitting a curve to model biological growth trend to each series, and dividing out the growth model. The chronology was computed as a robust estimation of the mean value function to remove effects of endogenous stand disturbances; this enhances the common signal contained in the data. In the 'residual' chronology, autoregressive modelling was performed on the detrended ring measurement series. Robust estimation of the mean value function produces a chronology with a strong common signal and without persistence. The pattern of variation is generally similar among trees growing throughout the same region for the same time period because of variations in macro climatic factors, which are common to these trees¹³. Synchronization of growth pattern of Blue Pine in most of these sites covering a wide area of the Ziro Valley suggests a common force controlling the growth in the region. Generally, common growth pattern is exhibited in trees where climate has a significant role in limiting the tree growth. The averaging of standardized ring-width values or indices

Table 2. Selected statistics of tree-ring site chronologies of *Pinus wallichiana* at Ziro Valley

	DOB	HAR	HON	MIC	RAG
Time-span	1704–2000	1792–2000	1859–2000	1773–2000	1855–2000
Number of years	297	209	142	228	146
Tr (Cr)	4 (10)	17 (28)	3 (6)	15 (35)	6 (12)
MS	0.182	0.161	0.207	0.137	0.147
SD	0.204	0.172	0.208	0.139	0.160
AC (1)	0.352	0.335	0.314	0.233	0.420
Common period	1865–88	1918–89	1904–2000	1837–98	1913–84
Number of years	124	72	97	162	72
Tr (Cr)	3(6)	13(17)	3(5)	11(19)	6(8)
Mean r_{mt}	0.234	0.176	0.183	0.163	0.183
Mean r_{wt}	0.565	0.230	0.256	0.351	0.264
Mean r_{bt}	0.151	0.174	0.165	0.153	0.176
SNR	1.832	3.642	1.124	3.907	2.021
EPS	0.647	0.785	0.529	0.796	0.669
VARpc1 (%)	36.5	24.6	36.6	21.9	29.2

Tr (Cr), Number of trees (number of cores); MS, Mean sensitivity; SD, Standard deviation; AC (1), Lag 1 autocorrelation; Mean r_{mt} , Mean correlation among all radii; Mean r_{wt} , Mean correlation within tree; Mean r_{bt} , Mean correlation between trees; SNR, Signal-to-noise ratio; EPS, Expressed population signal; VARpc1, Percentage of variance explained by the first eigenvector.

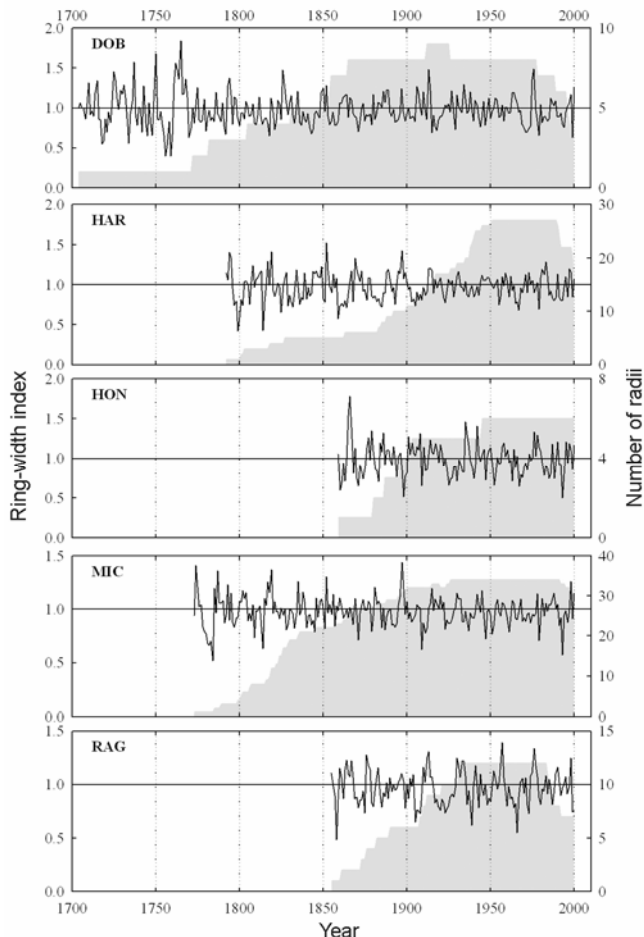


Figure 2. Five tree-ring site chronologies of *P. wallichiana* of Ziro Valley.

reduces the amount of variability due to non-climatic factors and enhances the ratio of climatic signal to non-

climatic signal in the mean chronology. Among the five site chronologies (Figure 2) of Blue Pine in the Ziro Valley, the longest one was from DOB which extends from AD 1704 to 2000, whereas the shortest one (AD 1855–2000) was from RAG. The statistical performance of five tree-ring site chronologies is shown in Table 2. The poor statistical performance of all the five tree-ring chronologies may be due to poor replication of the sample size and site disturbances due to human influence. The chronology which has a low autocorrelation, a high mean sensitivity and a high standard deviation has been considered suitable for dendroclimatic analysis¹³. In the study sites, the mean sensitivity and standard deviation were the highest (0.207 and 0.208 respectively) at HON and lowest (0.137 and 0.139 respectively) in MIC. Autocorrelation was highest (0.420) in RAG and lowest (0.233) in MIC. The mean correlation within the trees and among all the radii was highest (0.565 and 0.234 respectively) in DOB, whereas mean correlation between trees is highest in RAG (0.176) and the value was almost the same in HAR (0.174). In general, signal-to-noise ratio was low in these chronologies. It ranged from 1.124 to 3.907, in which HAR and MIC were slightly better in comparison to the other three chronologies. A value of 0.85 has been put forward as a reasonable threshold for expressed population signal (EPS)¹⁴, but none of these chronologies have crossed this value. The variances explained by HAR, MIC and RAG chronologies were much lower in comparison to the other two site chronologies, viz. DOB and HON (36.5 and 36.6% respectively).

Inter-site evaluations of the Blue Pine chronologies of Ziro Valley were pursued using three different statistical techniques: Pearson correlation matrix, principal component analysis (PCA) and hierarchical cluster analysis (HCA). These analyses help assess whether common variability exists within these site chronologies.

RESEARCH COMMUNICATIONS

The Pearson correlation shows that the three sites, HAR, HON and MIC have good correlation with each other (at $P = 0.01$), but their correlation with the other two sites, RAG and DOB is low though in between later these two sites have good correlation (0.233 at $P = 0.01$; Table 3). Except two, the correlation among other chronologies was moderate. This may be due to the number of samples incorporated in the site chronology. These two site chronologies (HAR and MIC) showing good correlation (0.446 at $P = 0.01$) have maximum number of samples compared to the other three sites, where the number of samples is less (Table 1).

PCA assesses the degree of similarity and site-related differences among chronologies. It is a data reduction or structure detection technique, and is widely used in dendroclimatology¹⁵. The original variables (residual tree-ring chronologies) have been transformed into a new set of uncorrelated variables (eigenvectors or principal components) in such a way that a minimum number of components explains a maximum percentage of the variance in the dataset. For the present study, PCA was computed on the correlation matrix of all five ring-width chronologies of Blue Pine covering a common period of AD 1859–2000. The number of components retained for further analyses was determined on the basis of the eigenvalue trace (or scree plot). Two principal components were identified with an eigenvalue >1.0 . These two components together explain 59.0% of the variance in the original dataset. Varimax rotation was used to aid in interpretation of the individual site loadings on each eigenvector¹⁶. Rotation of the principal components resulted in a clear separation of the sites (Table 4). It showed that 33.2% of the variance in tree-ring width was represented by HAR, HON and MIC in PC1 loading. Two sites, RAG and DOB fell in the PC2 loading, and both explained 25.8% of the total variance. Here RAG had the highest loading followed by DOB.

For HCA, Ward's method was employed using Pearson correlation coefficient as the measure of similarity. In the HCA dendrogram, two distinct groups (Figure 3) were identified which strongly agree with the natural orthogonal relationship noted in the correlation matrix (Table 3) and PCA (Table 4). Though it is not possible to establish the actual cause of this separation of the group, it seems to have resulted due to difference in the elevation of the tree-ring sites (low elevation 1245–1260 m and high elevation 1320–1582 m). However, a branch in the hierarchical tree suggests that these are from the same stock. This may be due to the year-to-year variability common in the major time-span between these two groups. This indicates that the signal change in between these chronologies seems to have resulted due to small-scale changes linked to the microclimatic conditions. This might be related to the distribution of trees in relation to elevation rather than distance between these sites or there may be other causes which are yet to be established.

Table 3. Pearson correlation of five tree-ring site chronologies of *P. wallichiana* at Ziro Valley

	DOB	HAR	HON	MIC	RAG
DOB	1				
HAR	0.146	1			
HON	0.167*	0.293**	1		
MIC	0.212*	0.446**	0.241**	1	
RAG	0.233**	0.125	-0.059	0.154	1

*Correlation is significant at the 0.05 level.

**Correlation is significant at the 0.01 level.

Table 4. Summary of principal component analysis for the common period 1859–2000 of five tree-ring site chronologies of *P. wallichiana* from Ziro Valley

Site	Principal component	
	PC1	PC2
HAR	0.752*	0.178
HON	0.737*	-0.168
MIC	0.695*	0.304
RAG	-0.079	0.860*
DOB	0.250	0.633*

*Rotated principal component with the highest loading for each chronology.

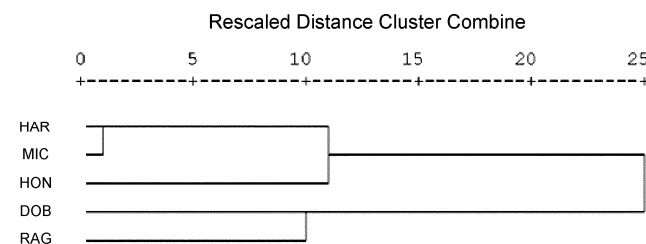


Figure 3. Dendrogram of *P. wallichiana* chronologies from hierarchical cluster analysis.

To evaluate the tree growth–climate relationship for a region, meteorological data (temperature and precipitation) from stations close to the sampling site are generally preferred, which could provide better understanding of climatic response to tree growth. These meteorological records are selected on the basis of their length, homogeneity and completeness of data. In the Northeast Himalayan region, stations having long climate records are not only limited in number, but most of them have high percentage of missing values thereby restricting their utility in qualitative dendroclimatic analyses.

The meteorological station, Ziro (93°42'E, 27°38'N, 1476 m asl), nearest to the sampling sites, has temperature and precipitation data extending from AD 1901 to 1980. Though it has long records, data for all the months during the years from 1963 to 1968 and 1977 are completely missing. Besides, more than six months' data are missing for the years 1901, 1962, 1971 and 1975. Overall for the period 1901–1980, the percentage of missing values in precipitation, minimum temperature and maximum tempe-

perature is 13.75, 14.58 and 13.65% respectively. In addition, there is a marked change in the quality of data; the variation in temperature was quite low in 1901–61 in comparison to 1969–80. This type of drastic change in climatic records generally happens when there is a shift of meteorological station from one site to other; but there is no such information for this station. Due to such discrepancies in these records, we prefer interpolated temperature and precipitation data (AD 1901–2000) of the nearest grid points (93°75'E, 27°75'N and 93°75'E, 27°25'N) of the CRU-TS 2.0 dataset¹⁷. The correlation between the annual temperature of the two grid points for the common period AD 1901–2000 was 0.982 ($P = 0.01$), whereas for precipitation, it was 0.959 ($P = 0.01$). Similarly, correlations for temperature and precipitation between data from meteorological station Ziro and average of two grid points for the common period AD 1902–1961 were 0.914 ($P = 0.01$) and 0.590 ($P = 0.01$) respectively. However, after eliminating the outlier value of 1952 from the precipitation dataset of the Ziro meteorological station, the correlation for the period 1902–51 was higher (0.837, $P = 0.01$). Time series of climate datasets from the meteorological station Ziro and average of two CRU TS 2.0 grid points along with bivariate plots for correlation are shown in Figure 4.

To assess the tree growth–climatic relationship of Blue Pine, bootstrap response functions¹⁸ were performed between the principal component scores of tree-ring chronologies and the interpolated climatic data. The applicability of this technique has been established earlier by several workers in tree ring–climate modelling^{19–22}. In this study climatic variables are taken for a 14-month period, i.e. from September of the previous year to October of the current year and the time-span has been taken from AD 1902 to 2000.

The response function of PC1 exhibited a positive response with February and April precipitation of the current year. The relationship was also positive with precipitation during December of the previous year and February of the current year in PC2 (Figure 5). Further, to understand the seasonal influences of precipitation on tree growth, we combined monthly climate data of some months which were found significant in limiting the growth of Blue Pine in this region, viz. December–February (DJF), February–April (FMA), December–April (DJFMA) and July–August (JA) in the computation of response function. In this analysis, among all these climatic variables, precipitation during pre-monsoon months, i.e. DJFMA was found highly significant. It has been recorded that the monsoon precipitation (JA) does not play a significant role in limiting the tree growth of this region. This is possibly due to the fact that the monsoon in this region never reaches the critical point to limit tree growth. In central Himalaya, Dolpo (western Nepal), growth of pine tree has also been found significantly correlated to the local summer rainfall, but not to the regional Indian Summer

Monsoon Index⁷. Precipitation during pre-monsoon creates cool and wet conditions to recharge the soil moisture that can benefit the growth during the ensuing growing season, as it help the initiation of cambial activity and rapid development of early wood cells. In case of temperature, PC1 and PC2 have a direct relationship with the December and November respectively, of the previous year. This relationship suggests that significant photosynthesis can occur in the evergreen needles during warm winters, which is used as stored energy in the ensuing growing season⁵.

The present tree-ring study of *P. wallichiana* at Ziro Valley, indicates that this taxon is suitable for dendroclimatic study for its clear and datable tree-ring sequences and synchrony in growth pattern. The climate signal (SNR) in these site chronologies of pine is moderate to low, which might have resulted due to intra-site difference. Generally, human disturbances and other natural calamities which vary from site to site may dilute the macro climatic signal, which is common to all these sites. These sites are highly disturbed due to clear-felling of trees for fuel and construction of houses, faulty agricultural practice ('jhum technique'), etc. By means of statistical analysis, viz. correlation matrices, PCA and HCA, two ecological zones for distribution of Blue Pine have been identified. This seems to be related due to the differences in elevation or changes in other site characteristics. A significantly positive response recorded in between tree growth and pre-monsoon precipitation indicates the significance of tree-ring data of Blue Pine not only in the pre-monsoon climatic reconstruction, but also in understanding of monsoon dynamics. Several studies have indicated that the

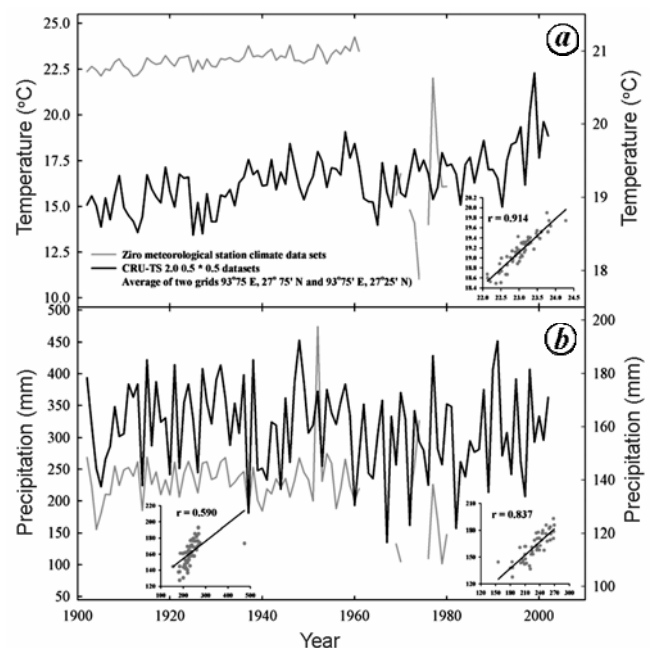


Figure 4. Comparison of climate data from meteorological station Ziro and interpolated CRU-TS 2.0 climate data. Bivariate plots for correlation.

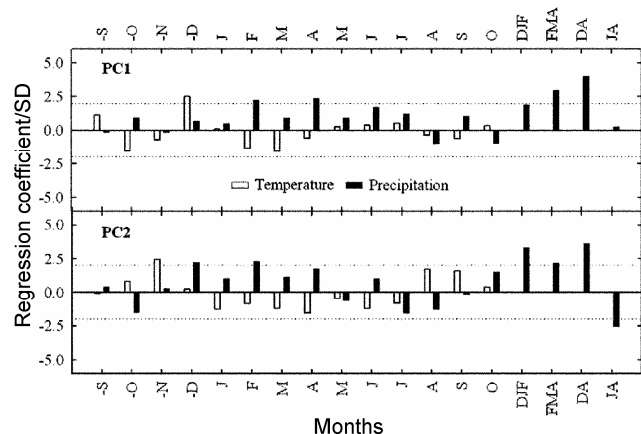


Figure 5. Response function plot of rotated principal component (PC) scores (PC1 and PC2) of *P. wallichiana* shown by regression coefficients divided by the standard deviation for the period 1902–2000. Dotted horizontal lines indicate significance level ($P < 0.05$) above and below the mean.

amount of precipitation in the form of snow during the previous winter and spring seasons regulates the build-up of the temperature gradient of the ensuing monsoon^{23–28}. Thus the dendroclimatic analysis from this region has a great promise for the analysis of the long-term linkage of the different climatic systems linked to monsoon in the global perspective. However, to obtain statistically meaningful and significant relationship we need long, good replication of tree-ring data from the less-disturbed sites of this region. There is an ample scope in the collection of tree-ring samples in the form of sub-fossil woods, buried under the peat and sediments in the river bed, in several places of Ziro Valley, which not only provide a scope for the replication of data but also to extend pine chronology through cross-dating.

- Sahni, K. C., In *Gymnosperms of India and Adjacent Countries*, Bishen Singh Mehendra Pal Singh, Dehradun, 1990, p. 169.
- Champion, H. G. and Seth, S. K., *A Revised Survey of the Forest Types of India*, Government of India Press, Delhi, 1968, p. 404.
- Bhattacharyya, A., LaMarche Jr, V. C. and Telewski, F. W., Dendrochronological reconnaissance of the conifers of northwest India. *Tree-Ring Bull.*, 1988, **48**, 21–30.
- Bhattacharyya, A., LaMarche Jr, V. C. and Hughes, M. K., Tree-ring chronologies from Nepal. *Tree-Ring Bull.*, 1992, **52**, 59–66.
- Yadav, R. R., Bhattacharyya, A. and Park, W.-K., Climate and growth relationship in Blue Pine (*Pinus wallichiana*) from the Western Himalaya, India. *Kor. J. Ecol.*, 1997, **20**, 95–102.
- Cook, E. R., Krusic, P. J. and Jones, P. D., Dendroclimatic signals in long tree-ring chronologies from the Himalayas of Nepal. *Int. J. Climatol.*, 2003, **23**, 707–732.
- Bräuning, A., Tree-ring studies in the Dolpo-Himalaya (western Nepal). In *TRACE, Tree Rings in Archaeology, Climatology and Ecology* (eds Jansma, E. et al.), Forschungszentrum Jülich, 2004, vol. 2, pp. 8–12.
- Bhattacharyya, A. and Yadav, R. R., Dendrochronological reconnaissance of *Pinus wallichiana* to study glacial behaviour in the western Himalaya. *Curr. Sci.*, 1996, **70**, 739–744.
- Singh, J. and Yadav, R. R., Tree-ring indications of recent glacier fluctuations in Gangotri, western Himalaya, India. *Curr. Sci.*, 2000, **79**, 1598–1601.

- Stokes, M. A. and Smiley, T. L., In *An Introduction to Tree-Ring Dating*, University of Chicago Press, Chicago, 1968, p. 73.
- Holmes, R. L., Computer-assisted quality control in tree-ring dating and measurement. *Tree-Ring Bull.*, 1983, **43**, 69–78.
- Grissino-Mayer, H. D., Evaluating crossdating accuracy: a manual and tutorial for the computer program COFECHA. *Tree-Ring Res.*, 2001, **57**, 205–221.
- Fritts, H. C., *Tree Rings and Climate*, Academic Press, New York, 1976, p. 567.
- Wigley, T. M. L., Briffa, K. R. and Jones, P. D., On the average value of correlated time series, with applications in dendroclimatology and hydrometeorology. *J. Climate Appl. Meteorol.*, 1984, **23**, 201–213.
- Briffa, K. R., Interpreting high-resolution proxy climate data – The example of dendroclimatology. In *Analysis of Climate Variability: Applications of Statistical Techniques* (eds von, H. and Navarra, A.), Springer Verlag, Berlin, 1995, pp. 77–94.
- Richman, M. B., Rotation of principal components. *J. Climatol.*, 1986, **6**, 293–335.
- Mitchell, T. D., Carter, T. R., Jones, P. D., Hulme, M. and New, M. A., Comprehensive set of high-resolution grids of monthly climate for Europe and the globe: the observed record (1901–2000) and 16 scenarios (2001–2100). Tyndall Centre Working Paper No. 55, 2004.
- Guiot, J., The bootstrapped response function. *Tree-Ring Bull.*, 1991, **51**, 39–41.
- Tardif, J., Brisson, J. and Bergeron, Y., Dendroclimatic analysis of *Acer saccharum*, *Fagus grandifolia*, and *Tsuga canadensis* from an old-growth forest, southwestern Quebec. *Can. J. For. Res.*, 2001, **31**, 1491–1501.
- Tardif, J., Conciatori, F. and Bergeron, Y., Comparative analysis of the climatic response of seven boreal tree species from northwestern Québec, Canada. *Tree-Ring Res.*, 2001, **57**, 169–181.
- Girardin, M.-P. and Tardif, J., Sensitivity of tree growth to the atmospheric vertical profile in the Boreal Plains of Manitoba, Canada. *Can. J. For. Res.*, 2005, **35**, 48–64.
- Piovesan, G., Biondi, F., Bernabei, M., Di Filippo, A. and Schirone, B., Spatial and altitudinal bioclimatic zones of the Italian peninsula identified from a beech (*Fagus sylvatica* L.) tree-ring network. *Acta Oecolog.*, 2005, **27**, 197–210.
- Dey, B. and Bhanukumar, O. S. R. U., The Himalayan winter snow cover area and summer monsoon rainfall over India. *J. Geophys. Res.*, 1983, **88**, 5471–5474.
- Shukla, J., Interannual variability of monsoons, In *Monsoons* (eds Fein, J. S. and Stephens, P. L.), Wiley Interscience Hoboken, NJ, 1987, pp. 399–464.
- Barnett, T. P., Dumenil, L., Schlese, U., Roeckner, E. and Latif, M., The effect of Eurasian snow cover on regional and global climate variations. *J. Atmos. Sci.*, 1989, **46**, 661–685.
- Douville, H. and Royer, J. F., Sensitivity of the Asian summer monsoon to an anomalous Eurasian snow cover within the Meteor-France GCM. *Climate Dyn.*, 1996, **12**, 449–466.
- Overpeck, J., Anderson, D., Trumbore, S. and Prell, W., The southwest Indian monsoon over the last 18,000 years. *Climate Dyn.*, 1996, **12**, 213–225.
- Yang, S., ENSO-snow-monsoon associations and seasonal-interannual predictions. *Int. J. Climatol.*, 1996, **16**, 125–134.

ACKNOWLEDGEMENTS. We thank Dr N. C. Mehrotra, Director, BSIP, Lucknow for encouragement and permission to publish this work. We also thank the forest officials of Ziro Valley, Lower Subansiri District, Arunachal Pradesh for permission and for providing necessary facilities during collection of samples. We are also grateful to IMD for providing meteorological data. This work is financially supported by Department of Science and Technology, New Delhi.

Received 23 July 2007; revised accepted 23 December 2008